

# The SST pattern effect on OLR: the role of large-scale convective aggregation

Reported by: Heng Quan ([hengquan@princeton.edu](mailto:hengquan@princeton.edu))

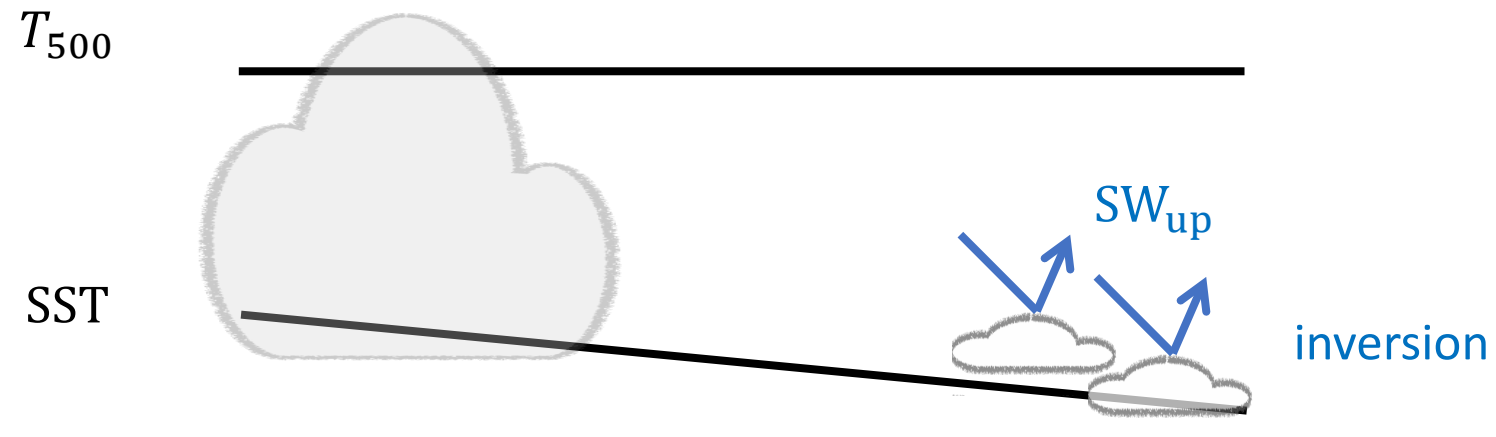
Princeton University, Atmospheric and Oceanic Sciences Program

Jun. 5<sup>th</sup>

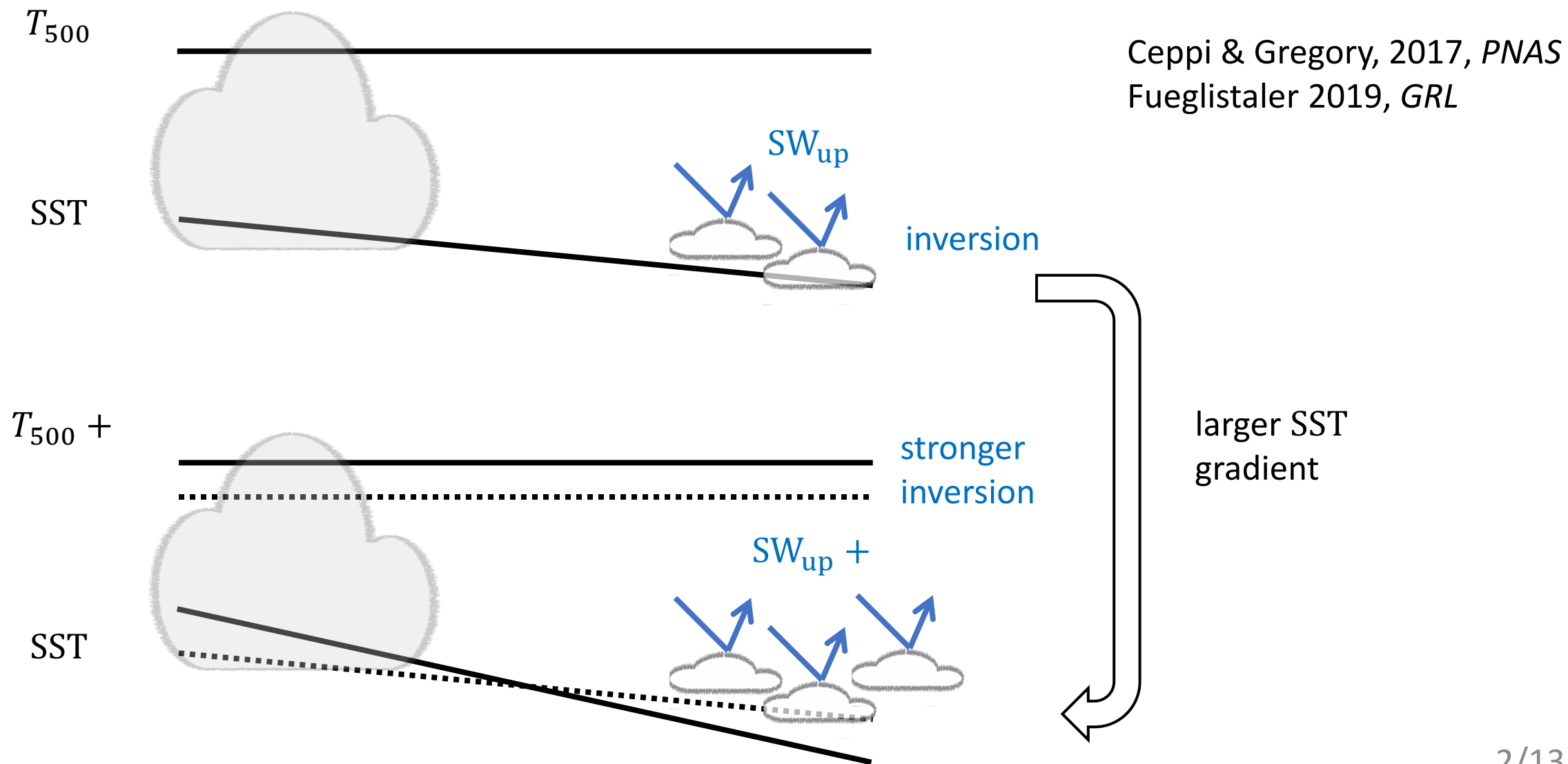
Collaborators: Stephan Fueglistaler, Bosong Zhang, Chenggong Wang



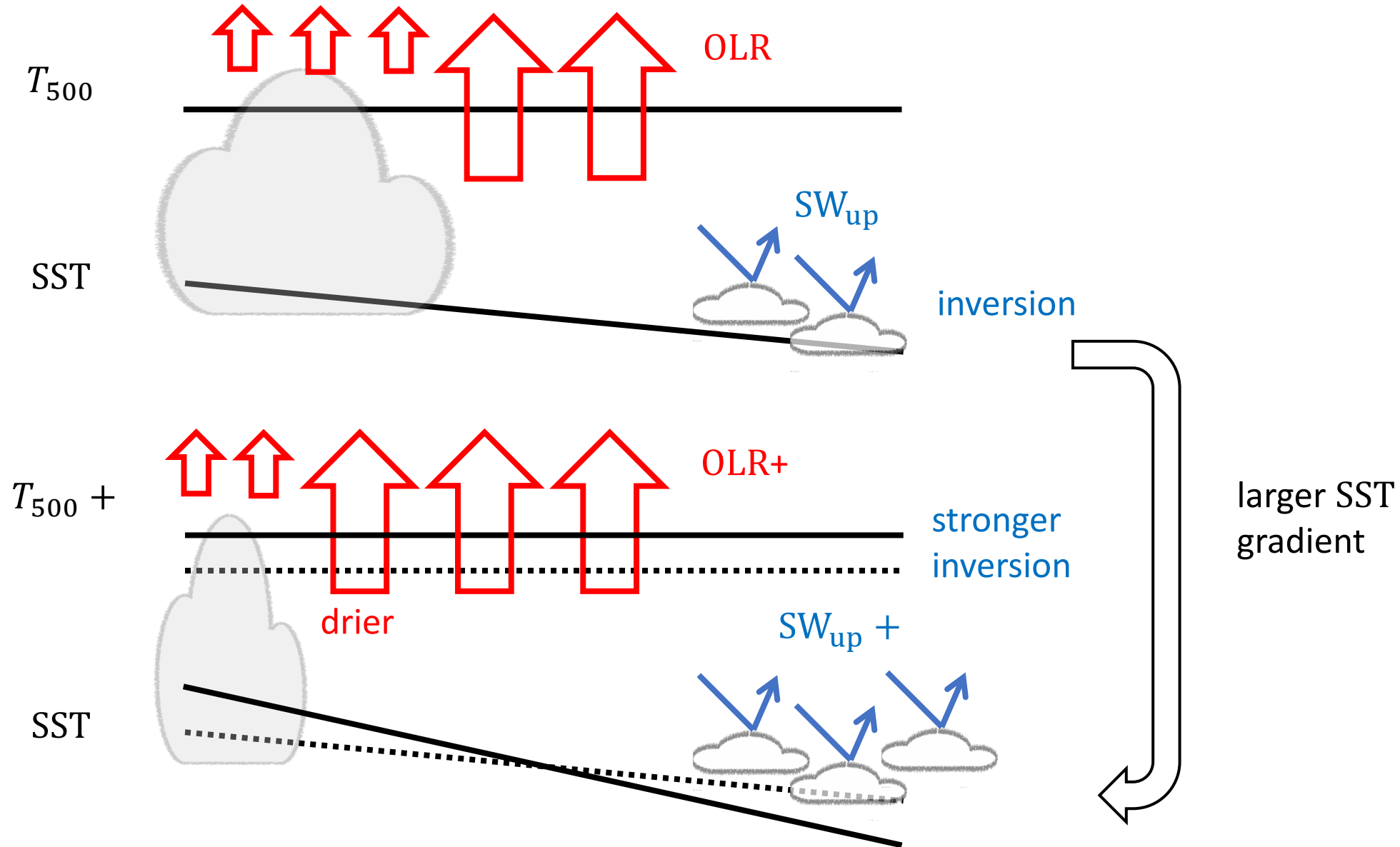
# The SST pattern effect – shortwave radiation



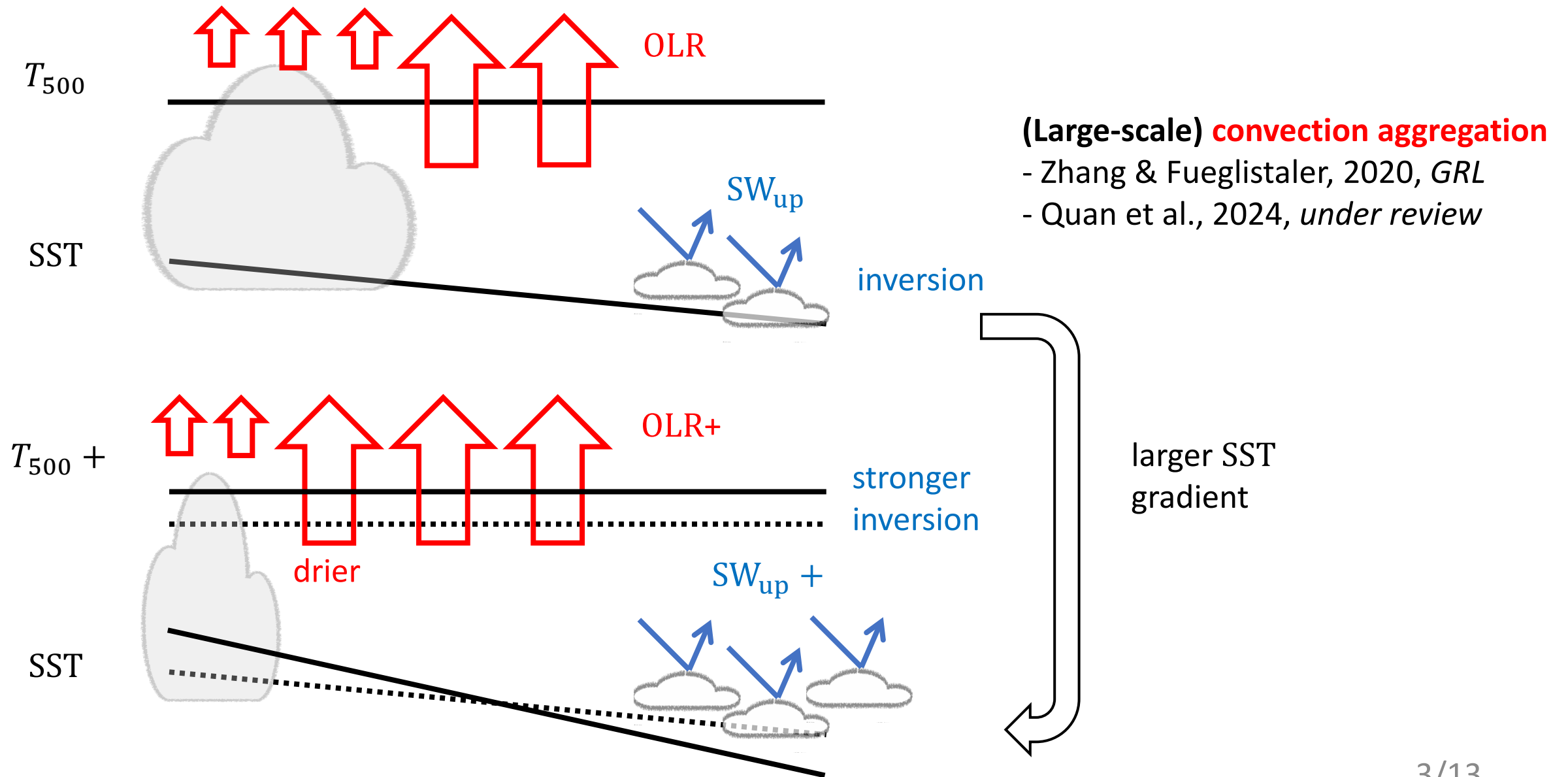
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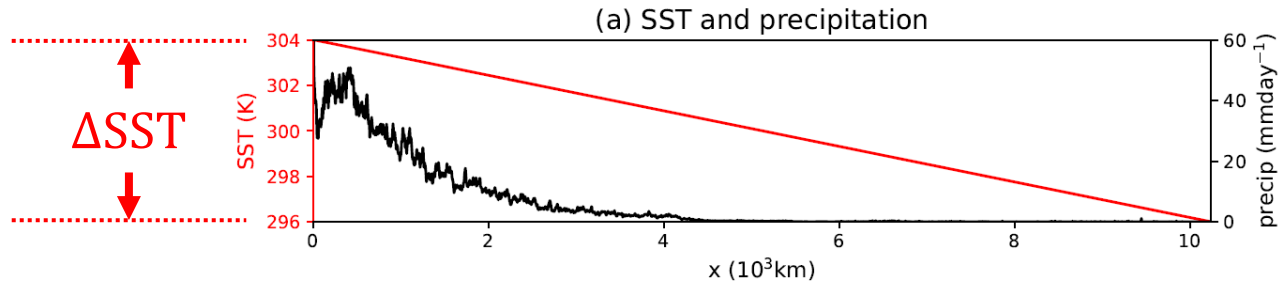
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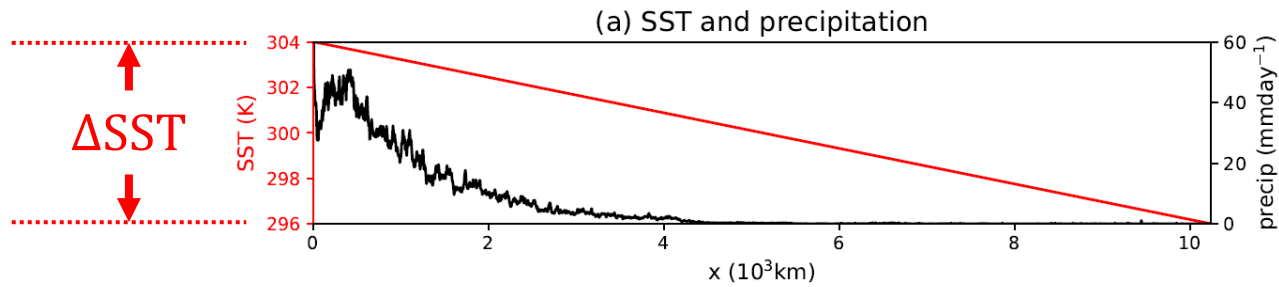


# Stronger convection aggregation $\rightarrow$ larger OLR: mechanism



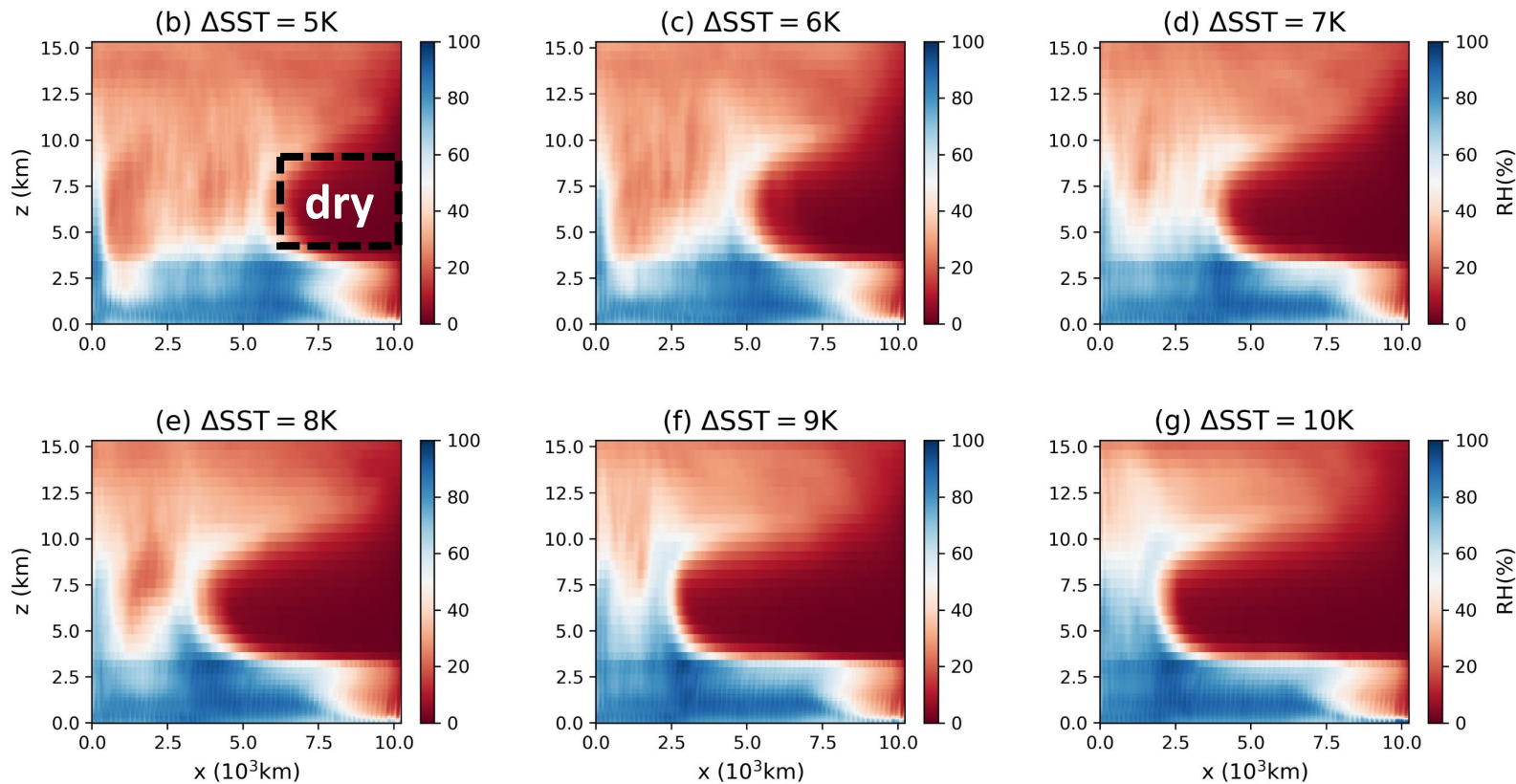
Cloud-resolving model - SAM  
2-D (x-z) mock Walker circulation  
Prescribed linear SST  
**Perturb  $\Delta\text{SST}$**

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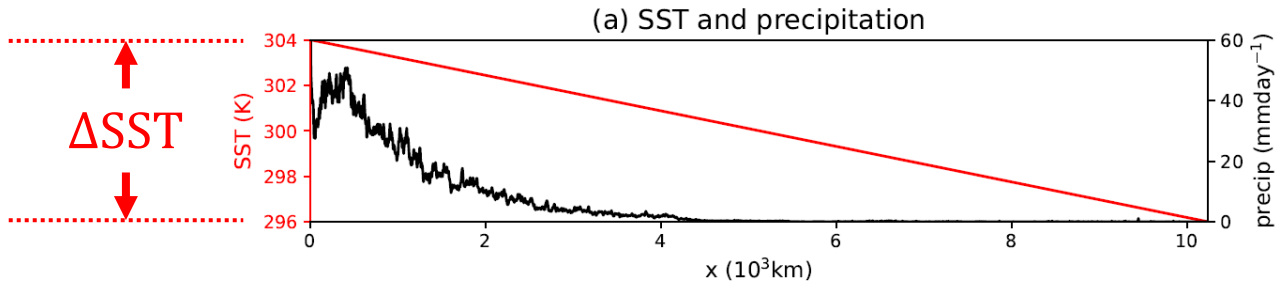


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**Perturb  $\Delta S_{ST}$**

RH(x, z) for different  $\Delta S_{ST}$

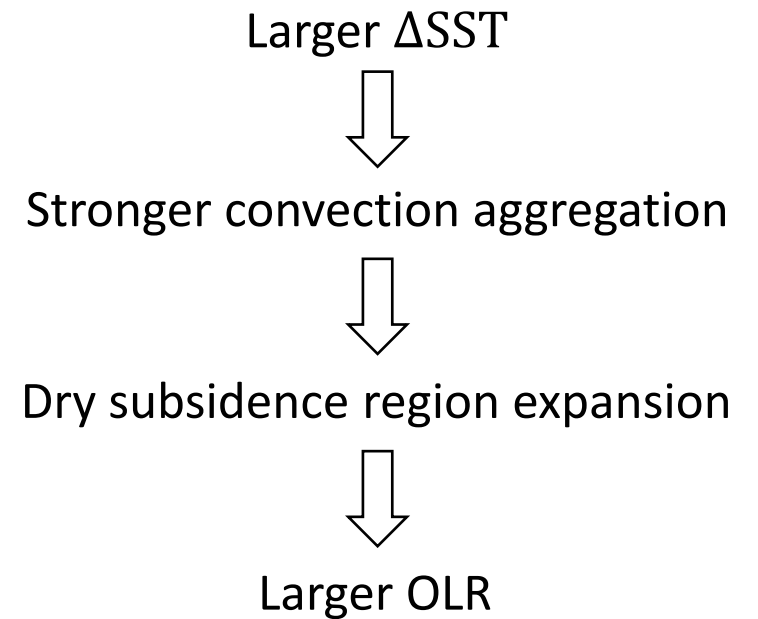
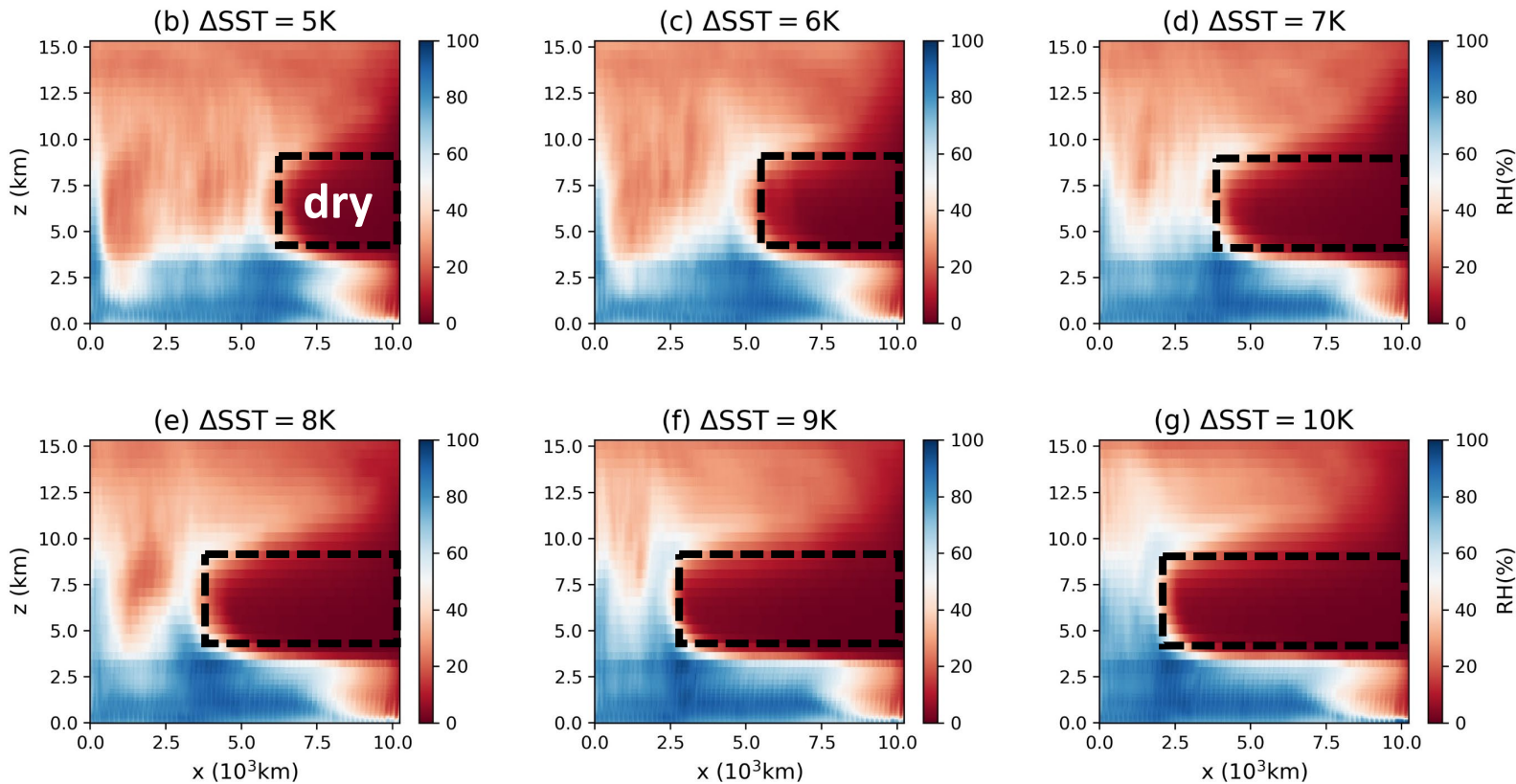


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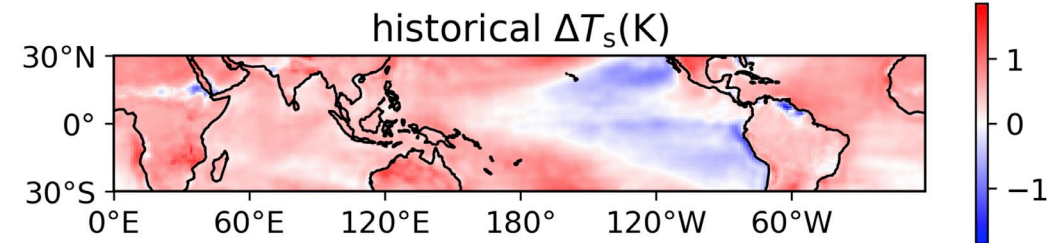
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1. Historical (1980-2010) SST pattern effect on OLR is comparable to reflected shortwave radiation because of convection aggregation strengthening.
2. The non-additivity error in the SST Green's Functions approach is explained by the non-additivity of convection aggregation.

# Historical SST pattern effect on OLR comparable to reflected SW

uniform: GFDL-AM4, uniform SST+4K

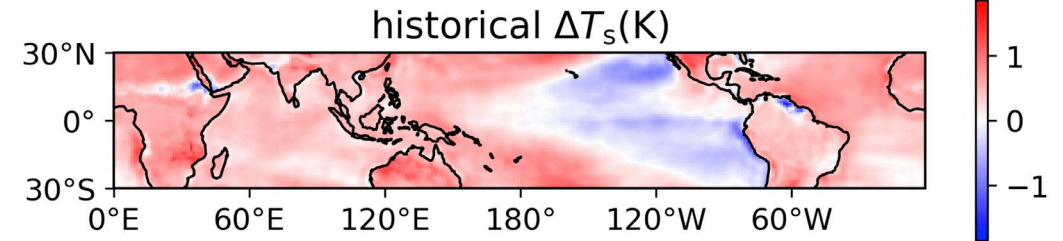
historical: SST perturbation = 1980 - 2010 SST trend  $\times$  30yr



# Historical SST pattern effect on OLR comparable to reflected SW

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Tropical (30°S ~ 30°N ) average responses, normalized to 1K mean surface warming

	$\frac{dSW_{up}}{dT_s}$	$\frac{dOLR}{dT_s}$	$\frac{dT_{500}}{dT_s}$	$\frac{dLCC}{dT_s}$	$\frac{dCRH}{dT_s}$	$\frac{dHCC}{dT_s}$
uniform	-0.21	1.69	1.44	-0.46%/K	0.27%/K	-0.19%/K
historical	1.72	3.14	1.95	1.51%/K	-0.62%/K	-1.10%/K
difference	<b>+1.93</b>	<b>+1.45</b>	+0.55	+1.97%/K	-0.89%/K	-0.91%/K

(historical – uniform)

500hPa T

Low cloud

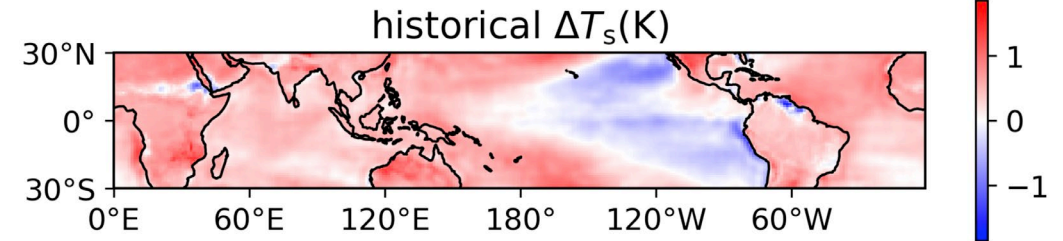
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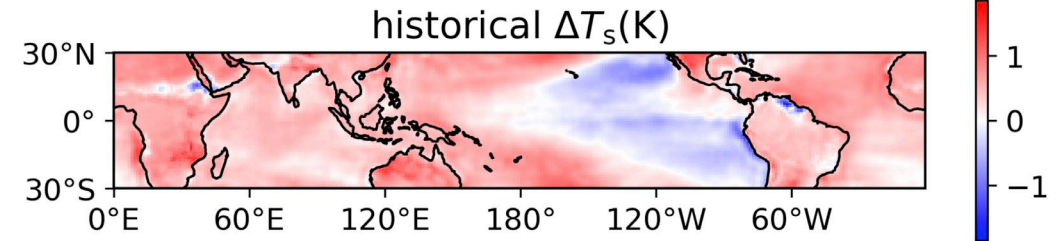
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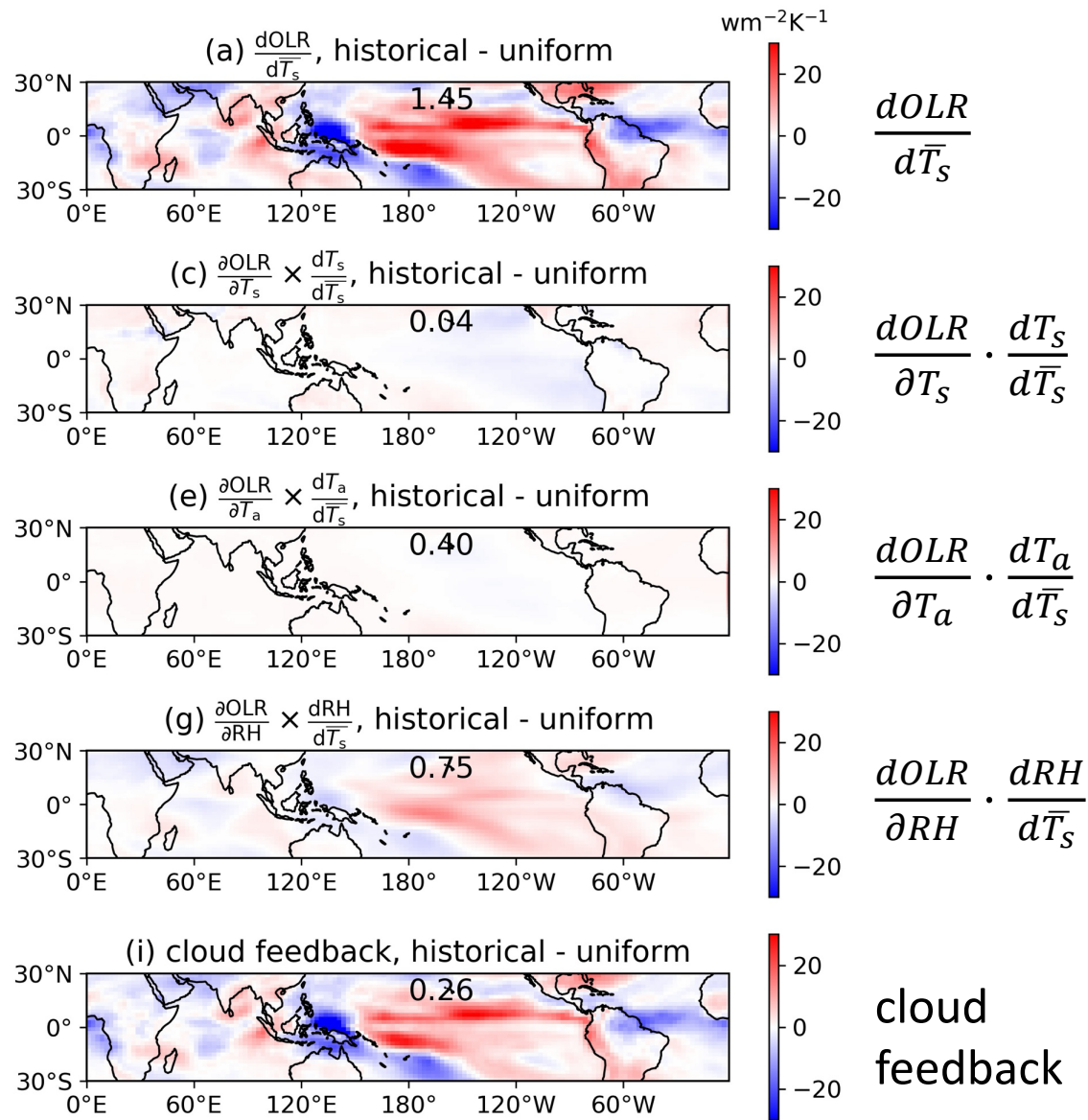
500hPa T

Low cloud

Column RH

High cloud

# Historical SST pattern effect on OLR due to convection aggregation



$$\frac{dOLR}{dT_s}$$

(Figures show historical – uniform)

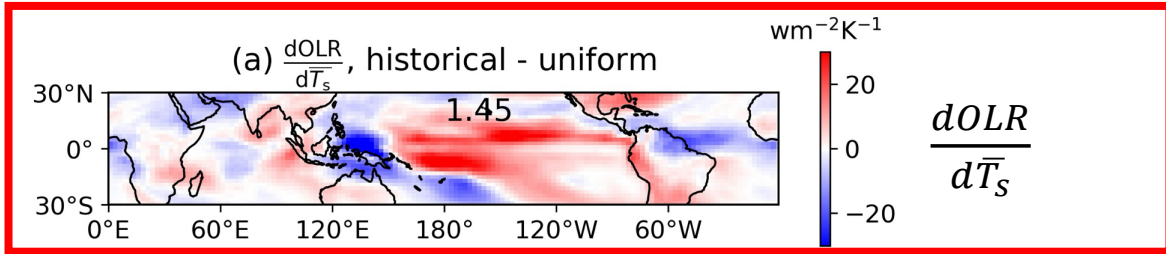
$$\frac{\partial OLR}{\partial T_s} \cdot \frac{dT_s}{d\bar{T}_s}$$

$$\frac{\partial OLR}{\partial T_a} \cdot \frac{dT_a}{d\bar{T}_s}$$

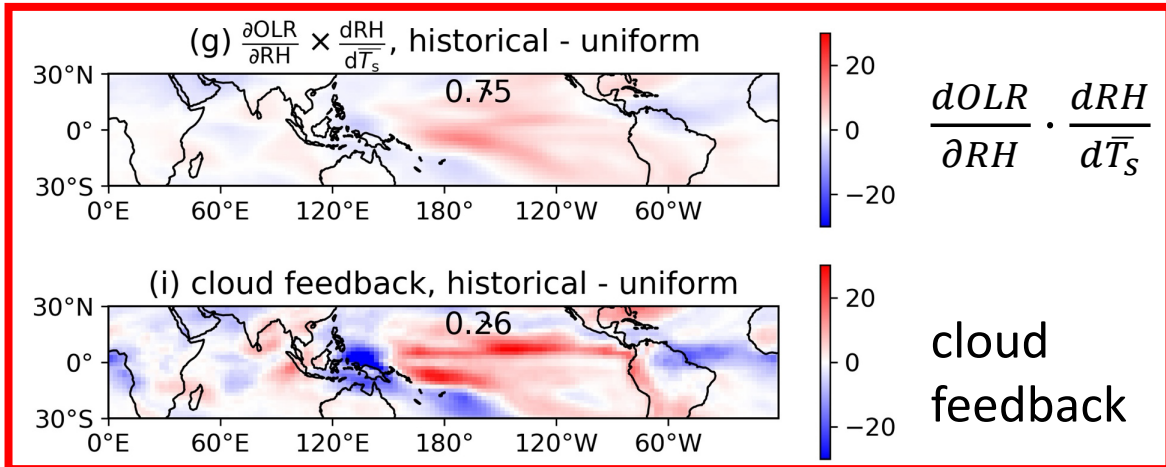
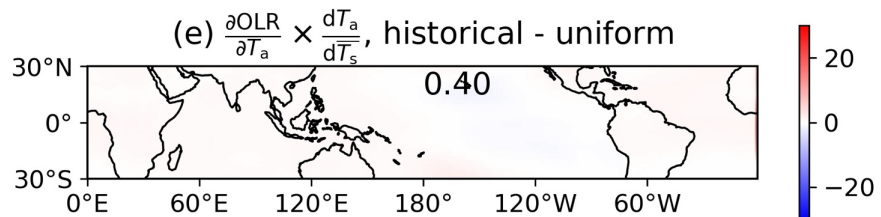
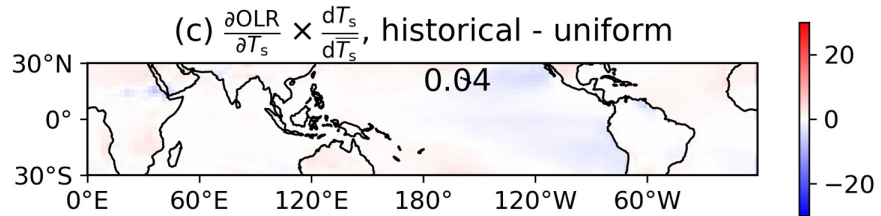
$$\frac{\partial OLR}{\partial RH} \cdot \frac{dRH}{d\bar{T}_s}$$

cloud  
feedback

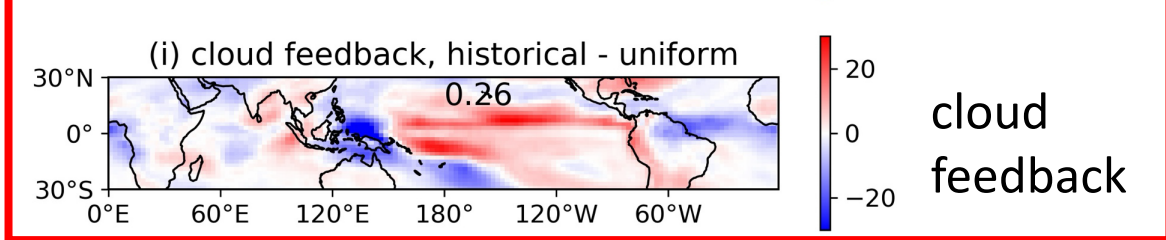
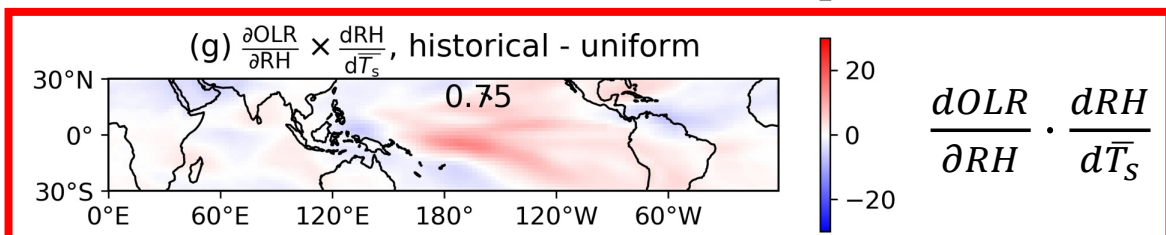
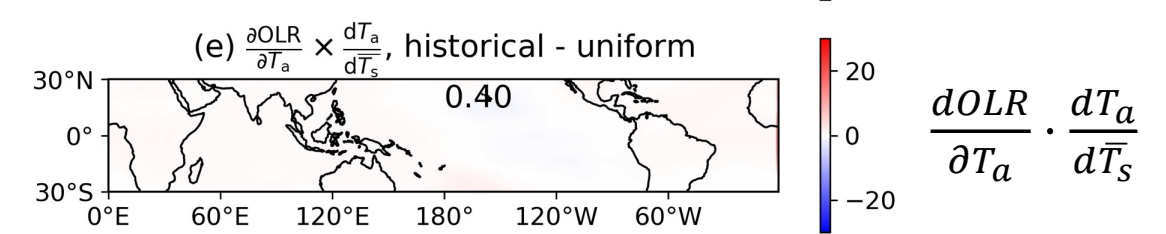
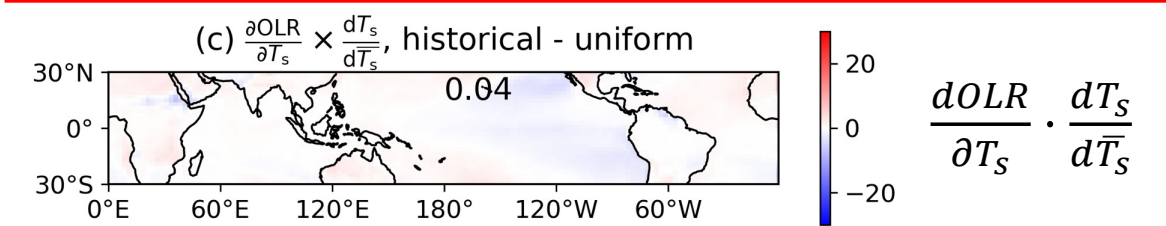
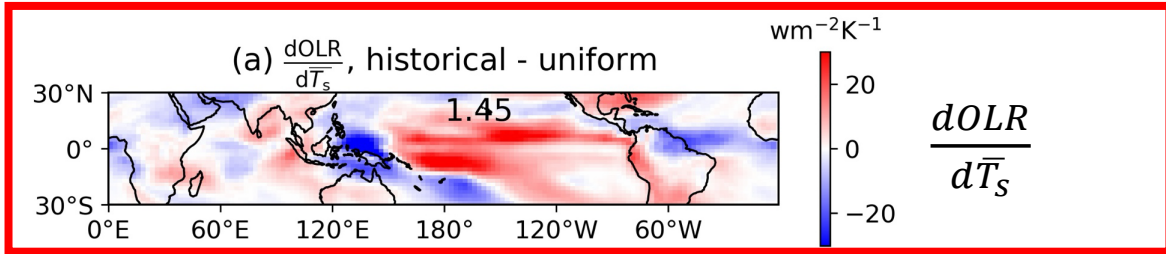
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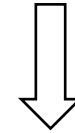
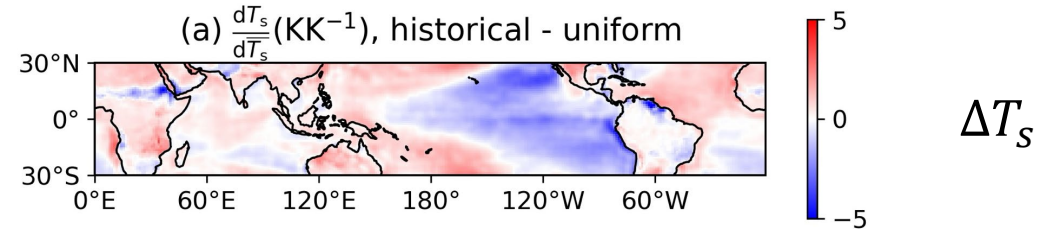
(Figures show historical – uniform)



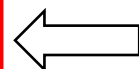
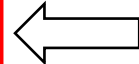
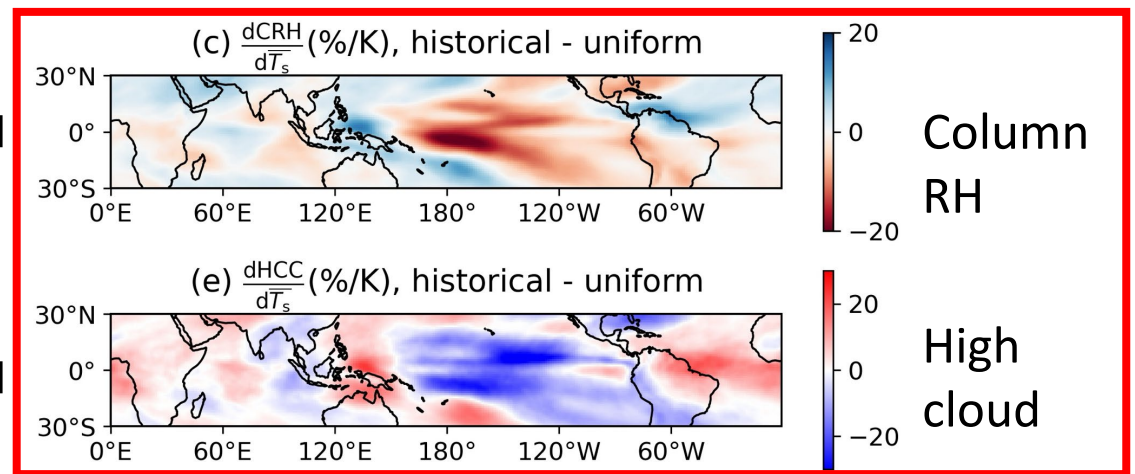
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(Figures show historical – uniform)



convection aggregation





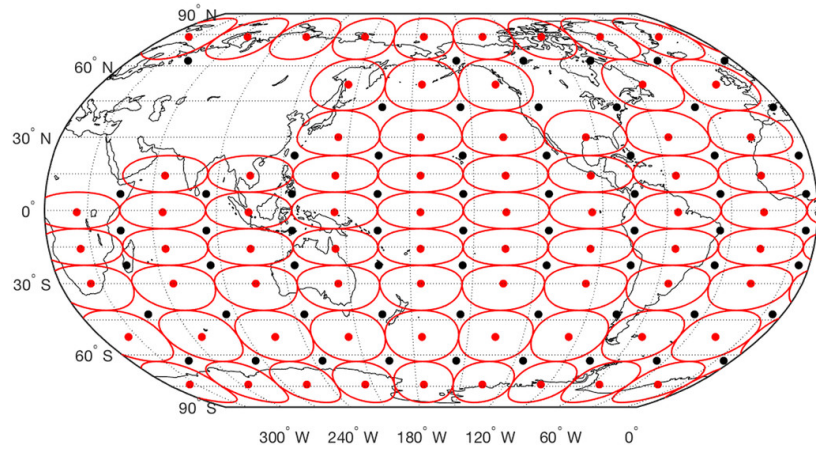
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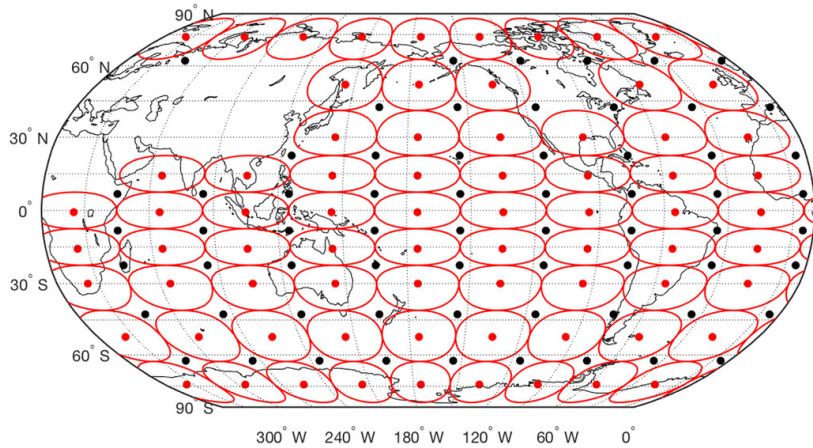
$$\Delta \bar{R} = \sum_j \frac{\partial \bar{R}}{\partial \text{SST}_j} \Delta \text{SST}_j$$



Dong 2019

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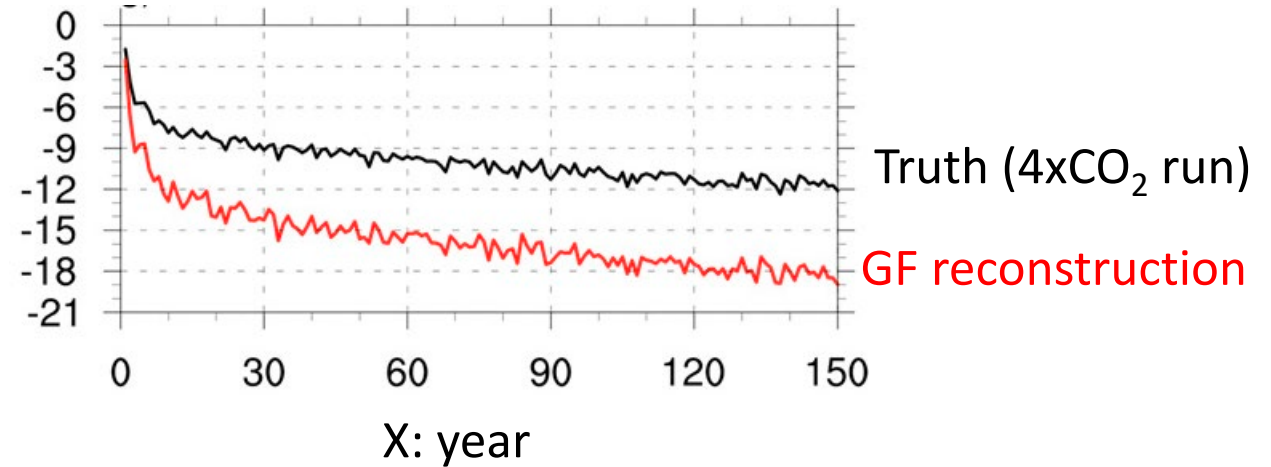
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Dong 2019

Y:  $\overline{\Delta R_{lw}}$

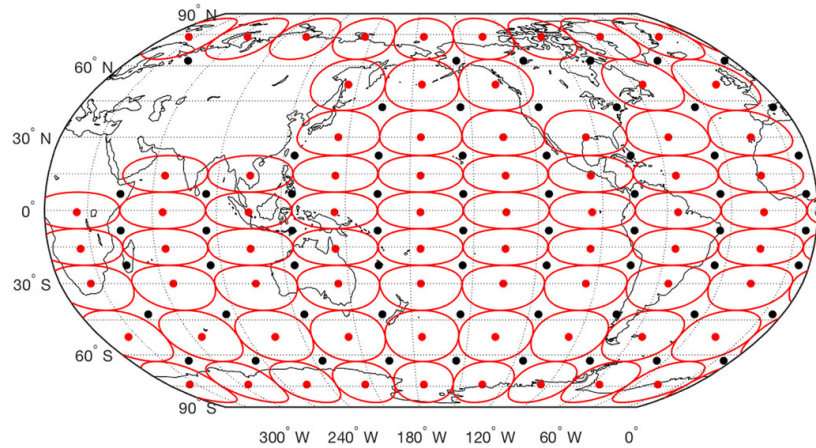
AGCM  $\leftarrow$  SST warming in 4xCO<sub>2</sub> GCM



Zhang et al., 2023, *JCLI*

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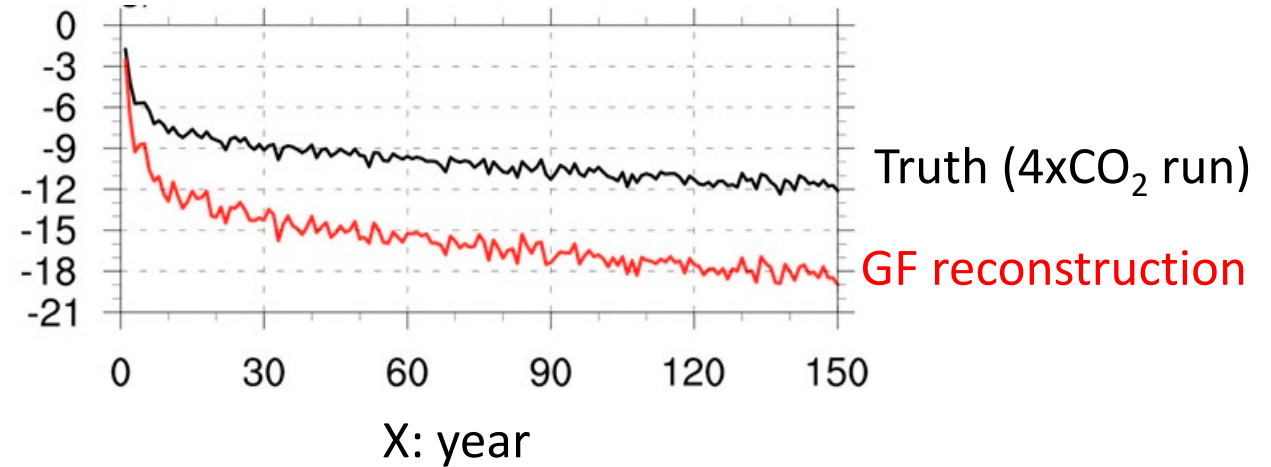
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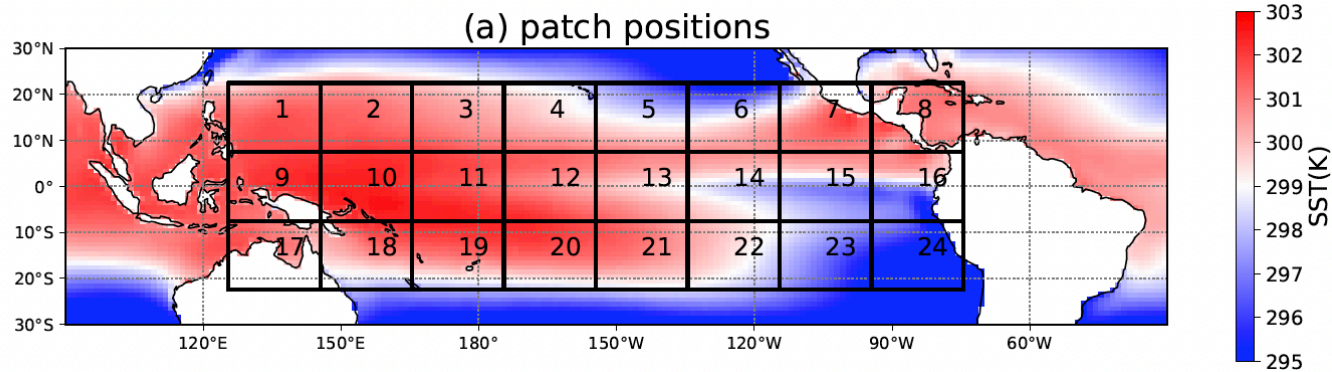


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$$\Delta \bar{R}(\Delta \text{SST}_1, \Delta \text{SST}_2, \dots) \neq \Delta \bar{R}(\Delta \text{SST}_1, 0, 0, \dots) + \Delta \bar{R}(0, \Delta \text{SST}_2, 0, \dots) + \dots$$

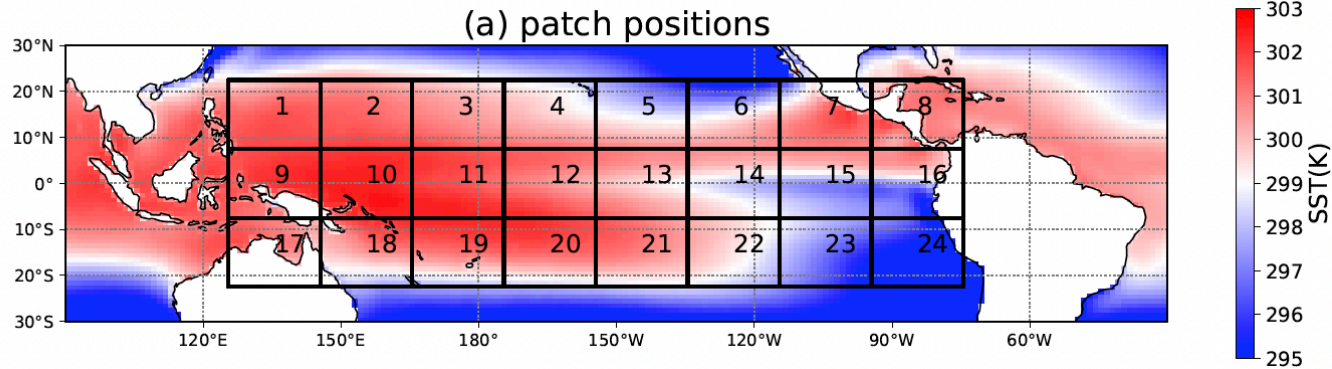
Question: Why is OLR response **overestimated** by linear sum?

# Two-patch combination: Linear sum overestimates OLR



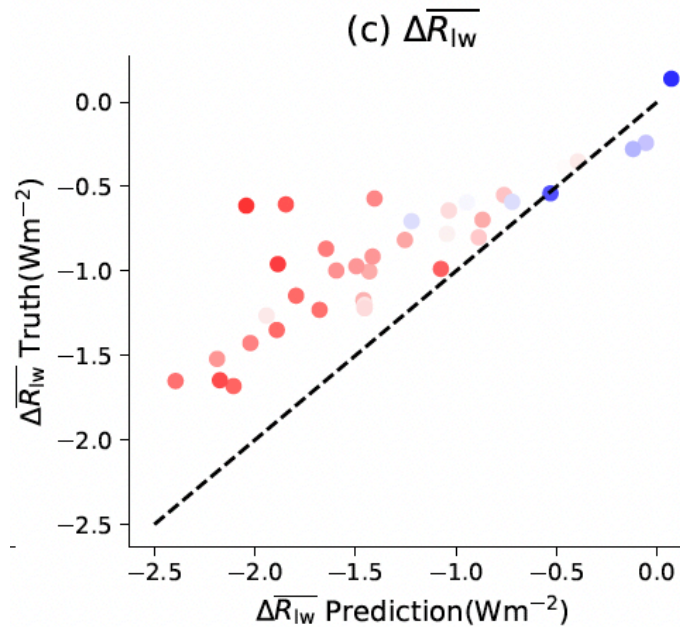
AM4, SST+4K perturbation in **two adjacent** tropical Pacific patches (37 combinations in total)

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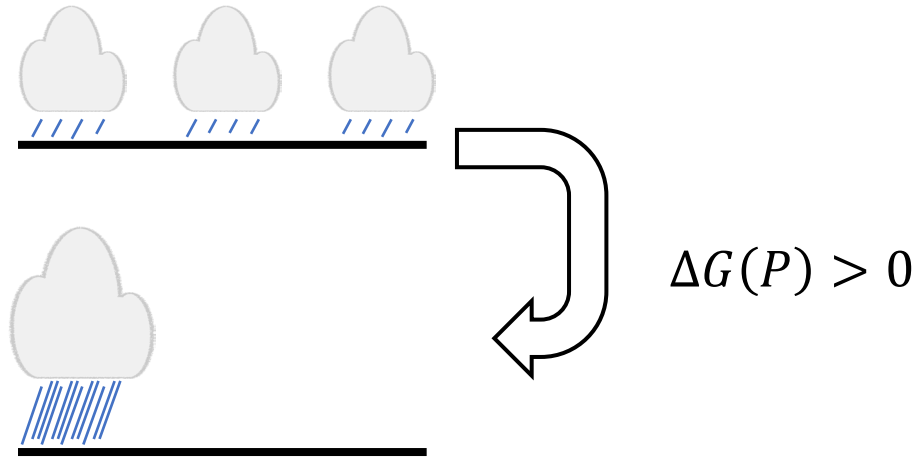
Y: actual response  
 $\Delta\bar{R}(\Delta SST_i, \Delta SST_j)$



Color: average SST of two patches

X: linear sum  
 $\Delta\bar{R}(\Delta SST_i, 0) + \Delta\bar{R}(0, \Delta SST_j)$

# OLR overestimation ← convection aggregation overestimation



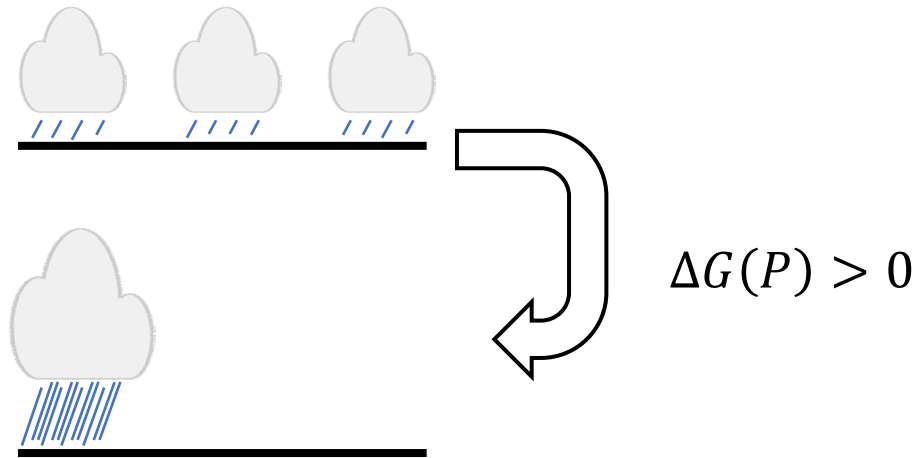
Tropical rainfall **Gini index**:  $0 < G(P) < 100$

If  $\Delta G(P) > 0$ :

- Precipitation more spatially uneven
- Convection more aggregated

- Zhang & Fueglistaler, 2020, *GRL*

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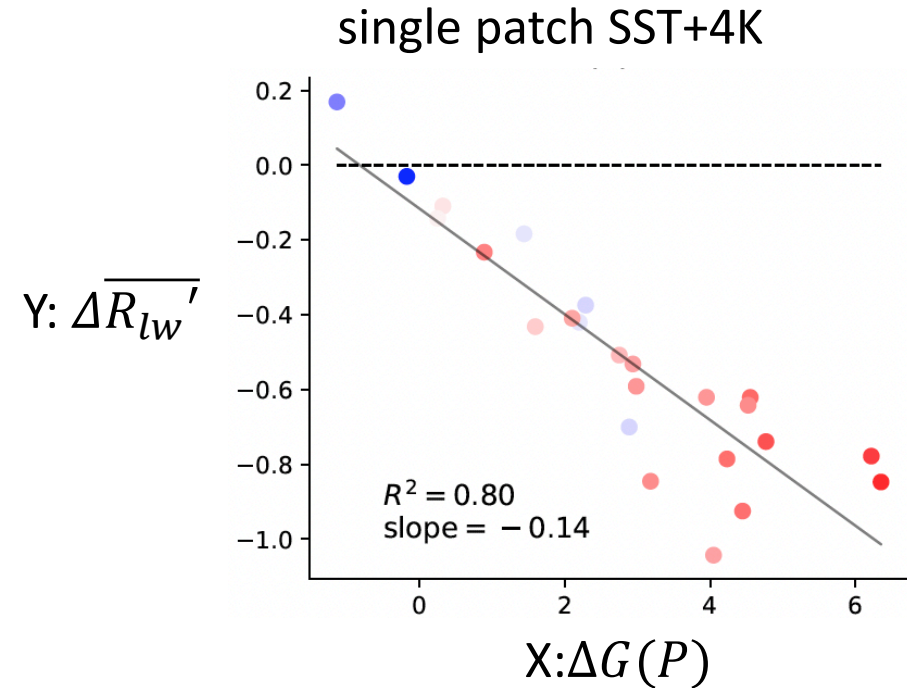


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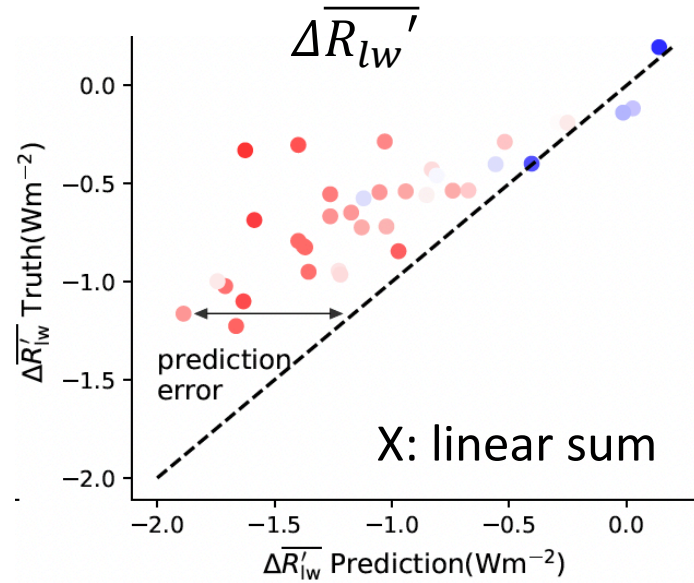


$$\Delta \overline{R_{lw}'} \propto \Delta G(P)$$

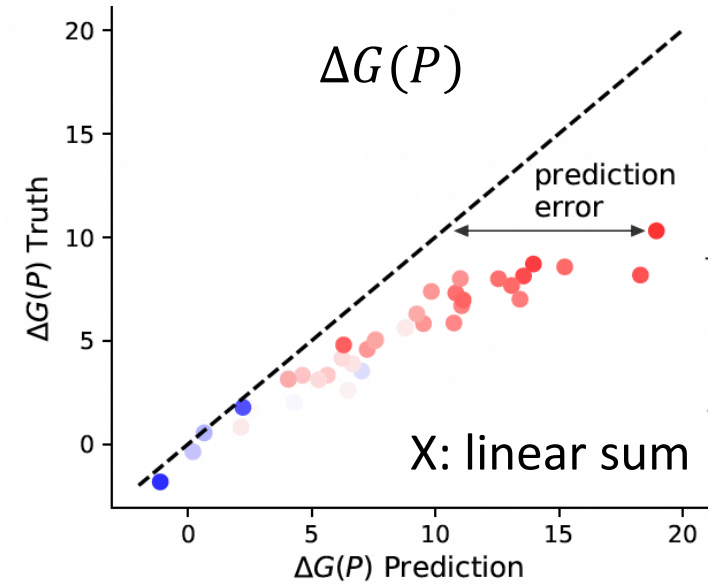


# OLR overestimation ← convection aggregation overestimation

Y: actual response



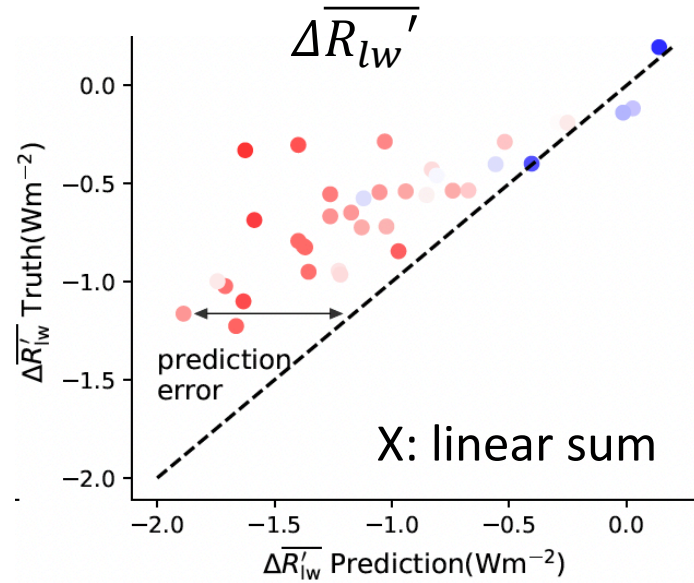
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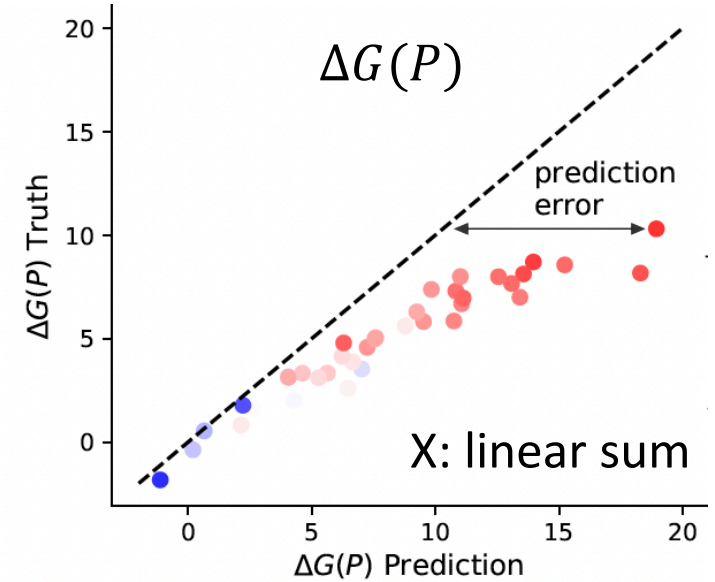
two adjacent patches SST+4K

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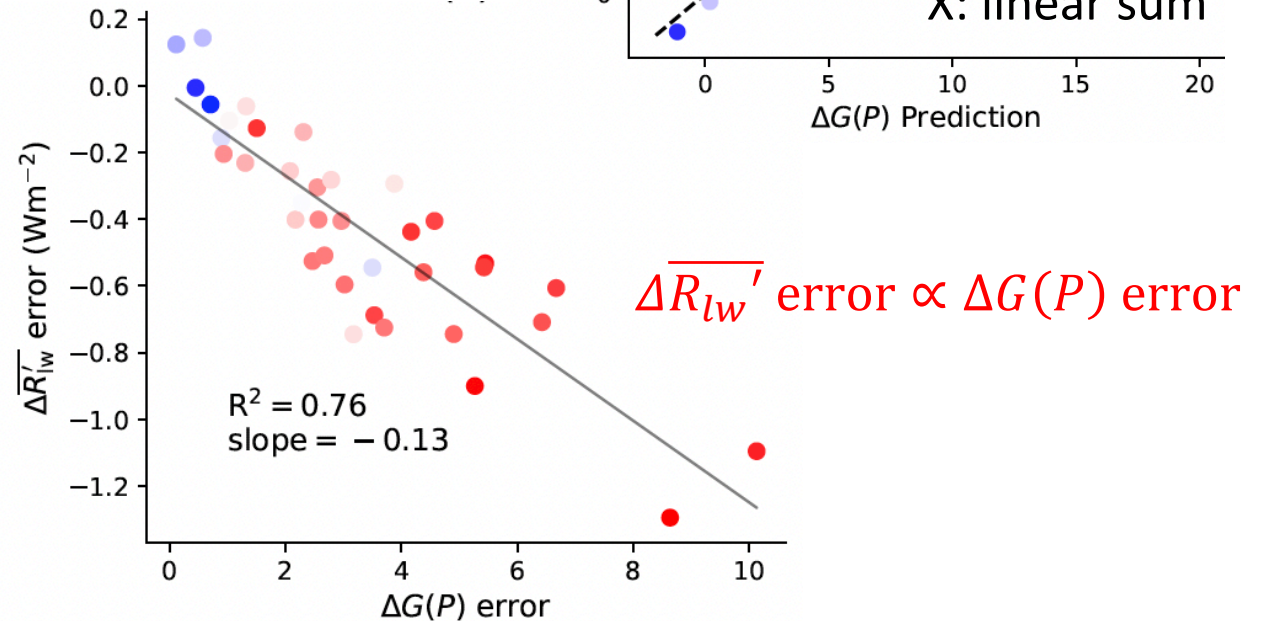
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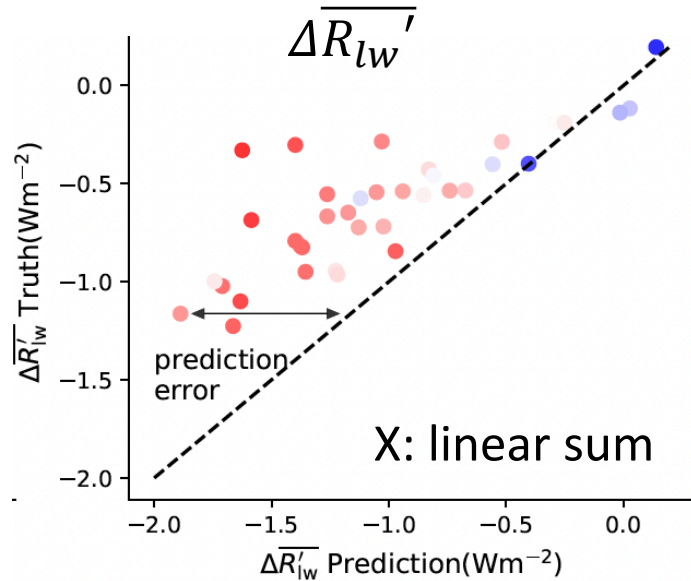


two adjacent patches SST+4K

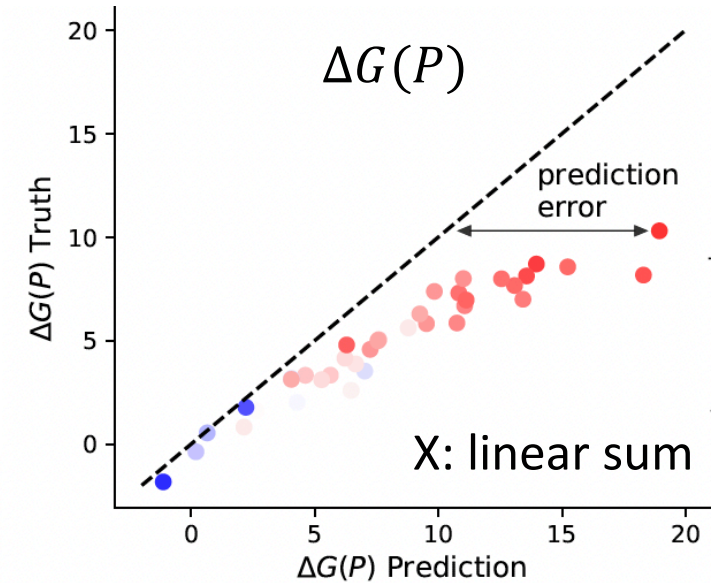


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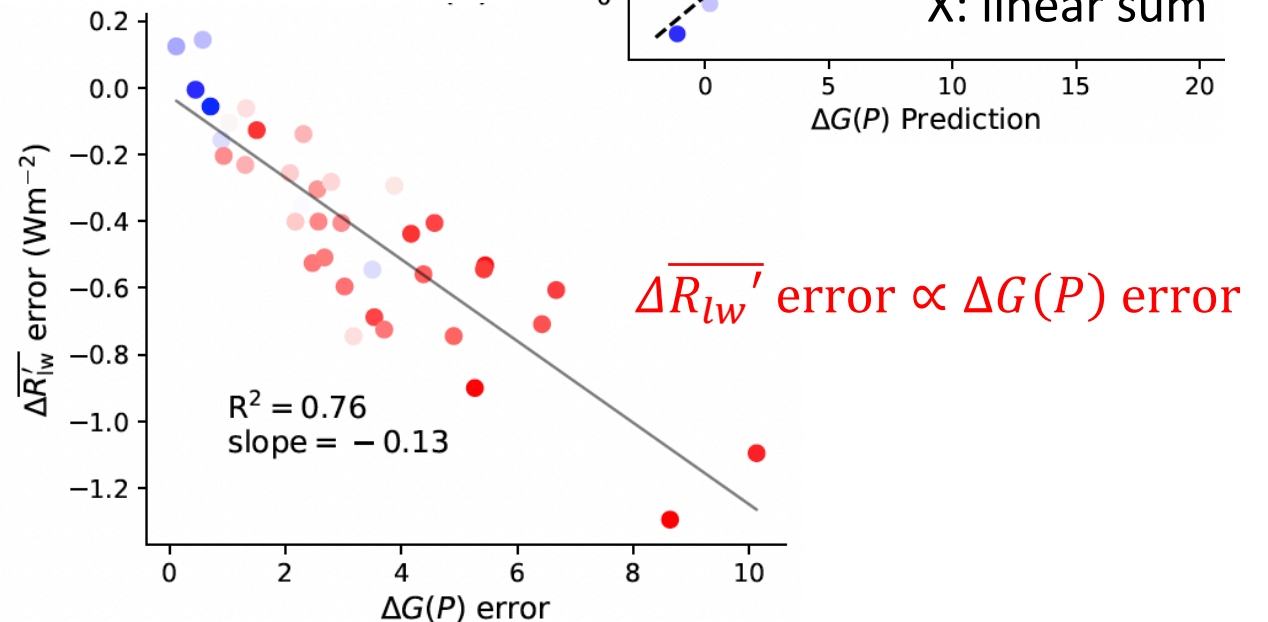
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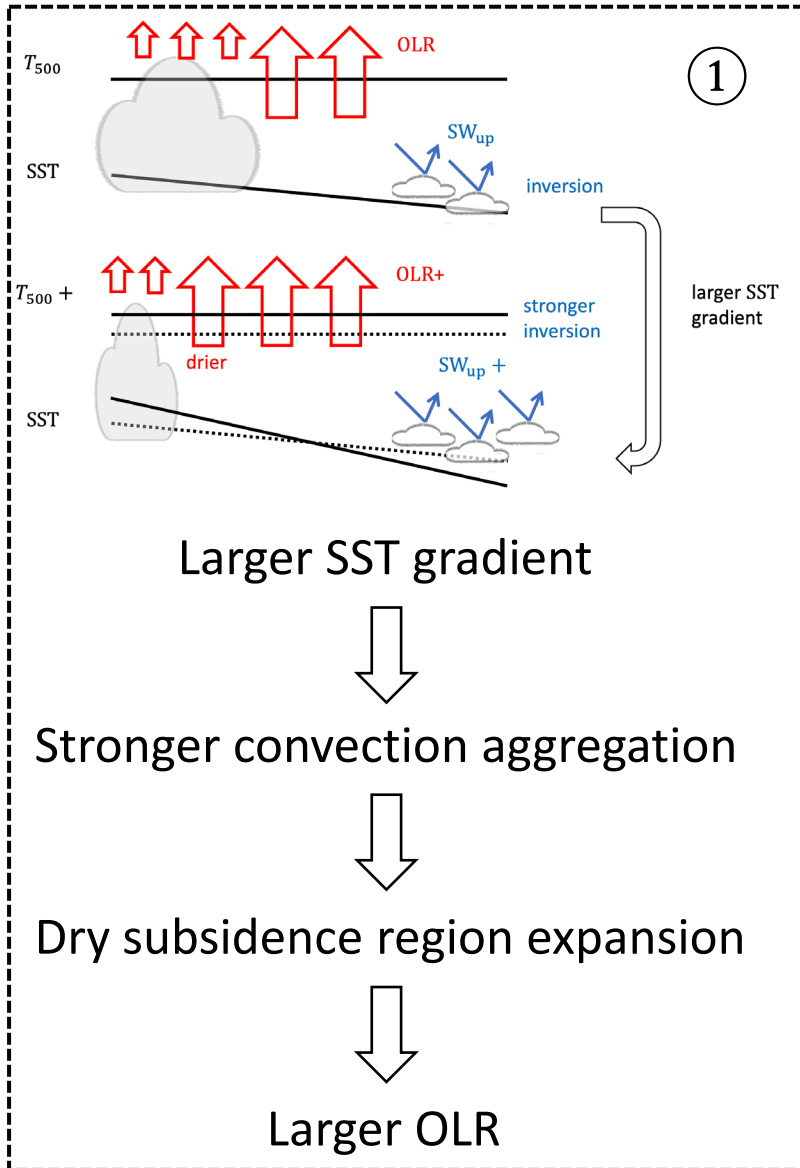
two adjacent patches SST+4K

Why is convection aggregation response overestimated by linear sum?

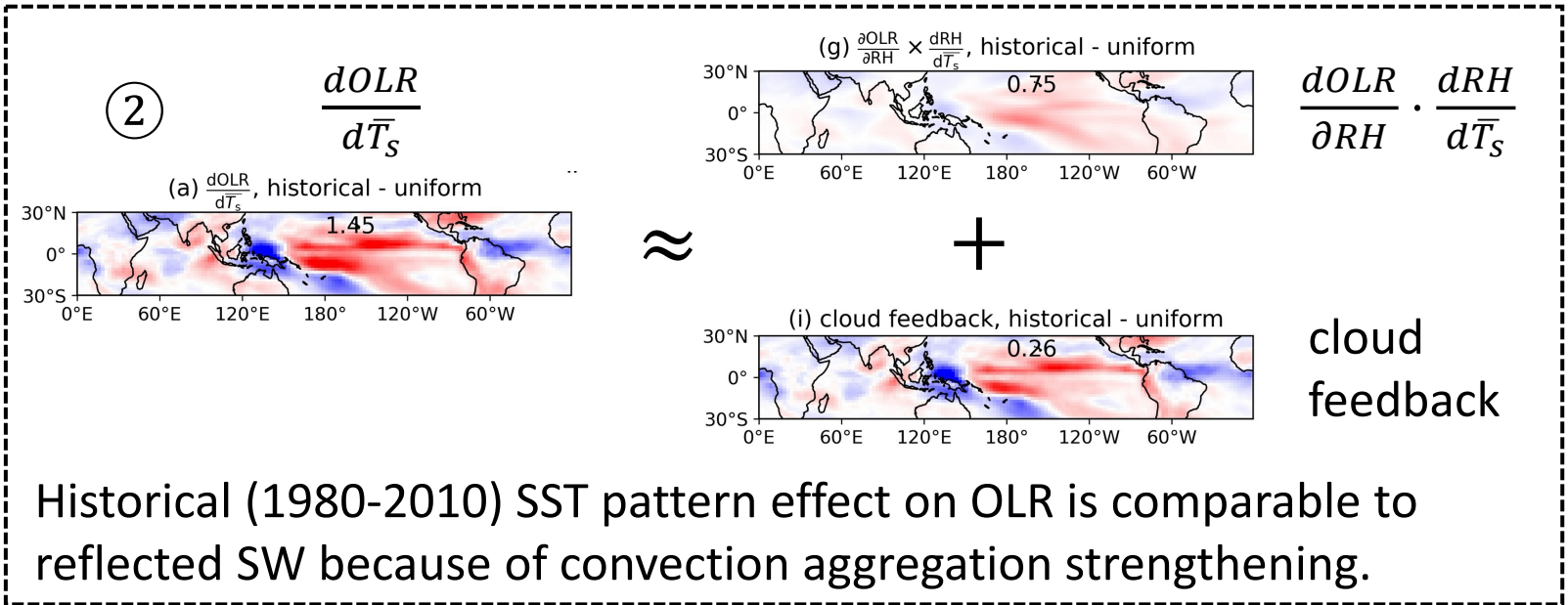
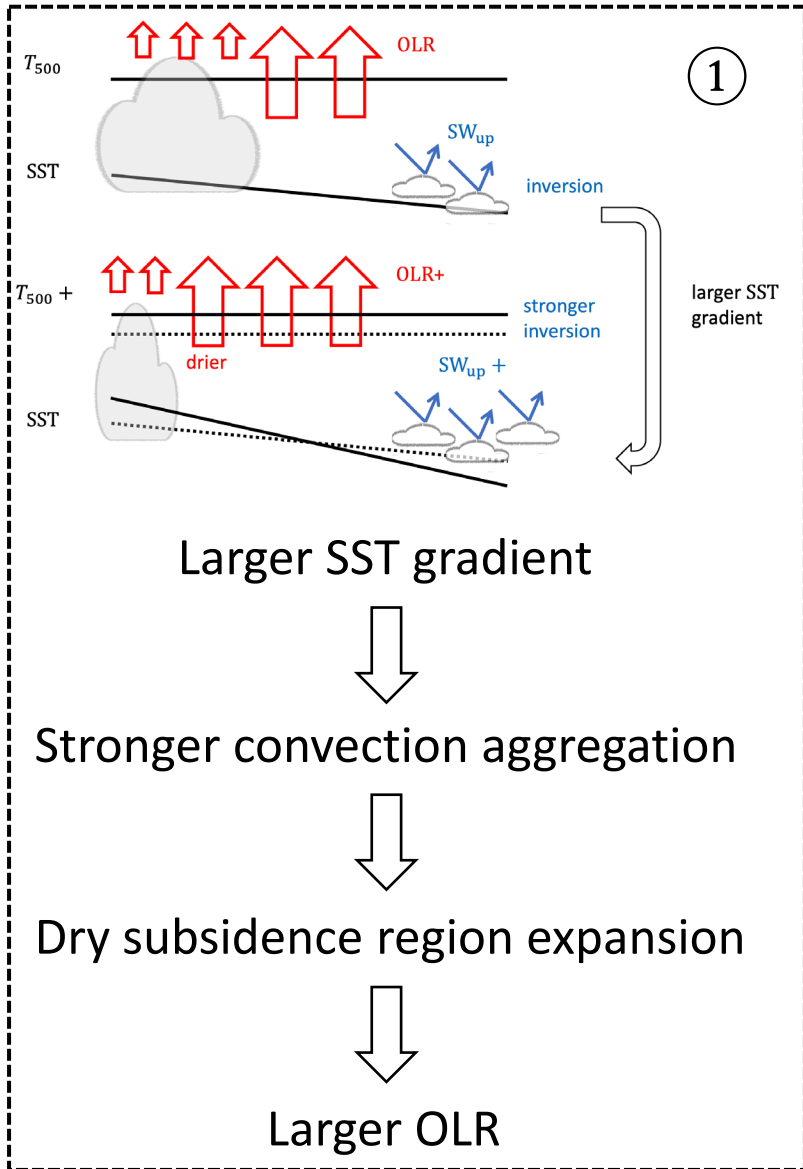
- non-additivity in circulation response
- Quan et al., 2024, *under review*



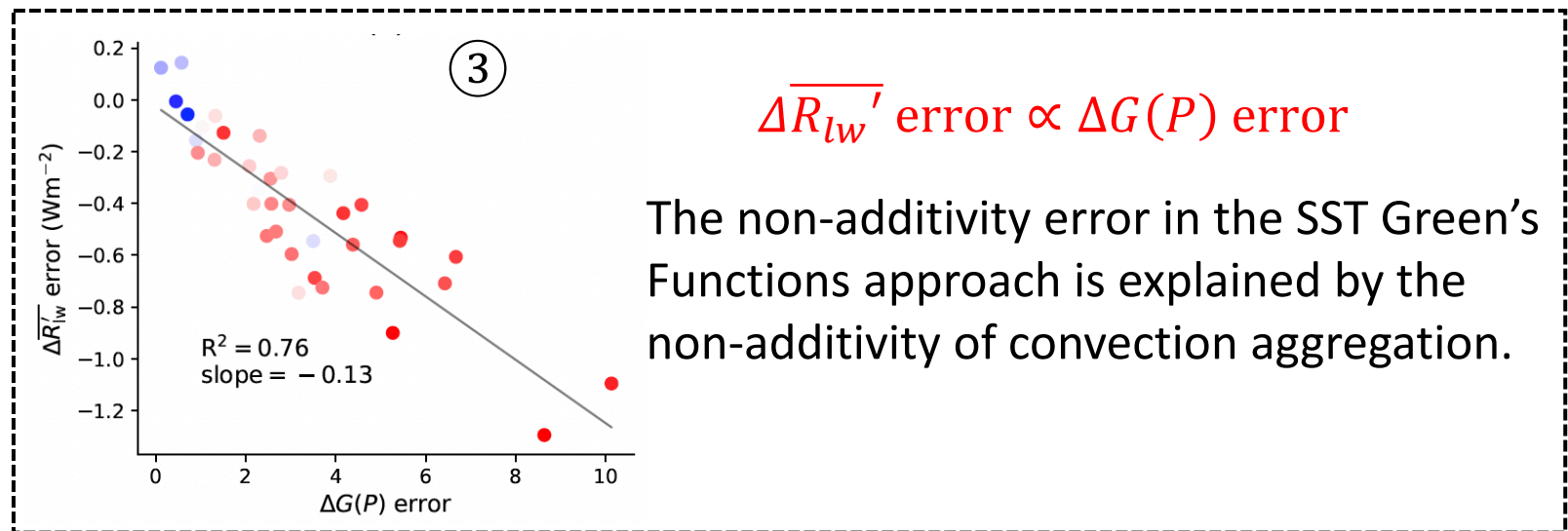
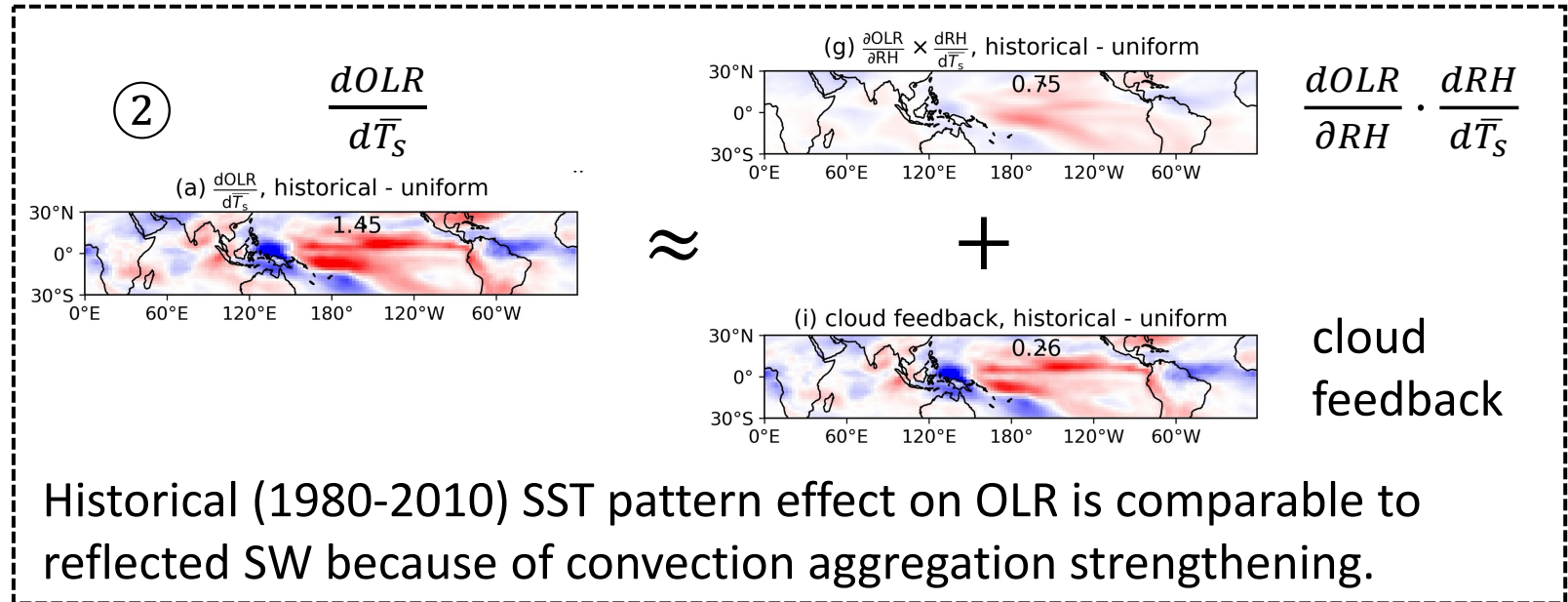
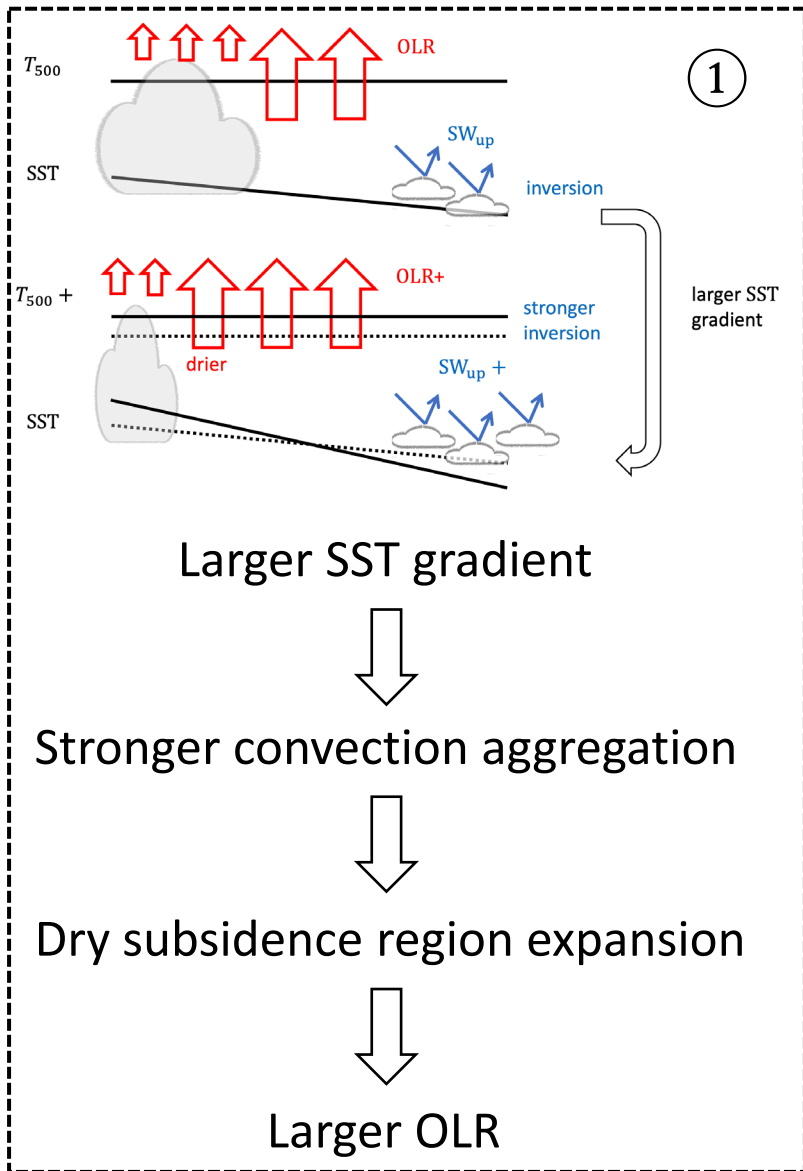
# Summary



# Summary



# Summary



Thanks!

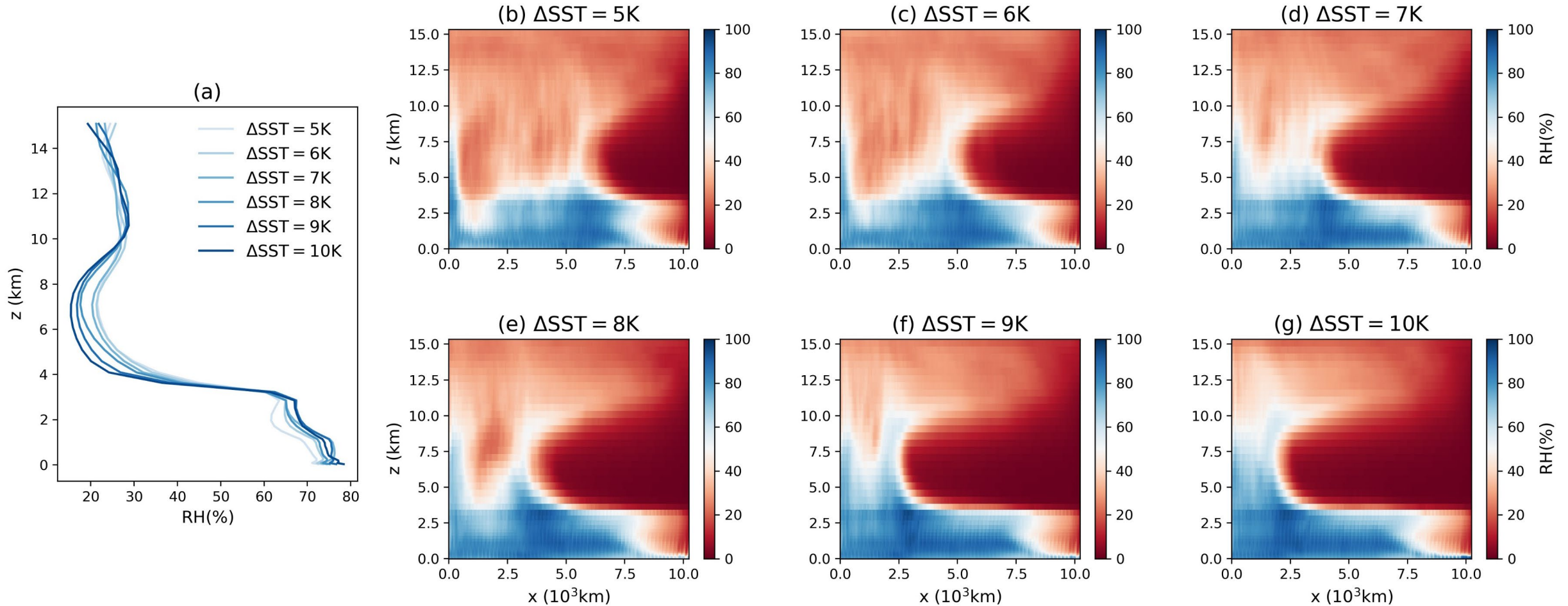
Preprint & Contact me:

[hengquan@princeton.edu](mailto:hengquan@princeton.edu)

<https://heng-quan.github.io>



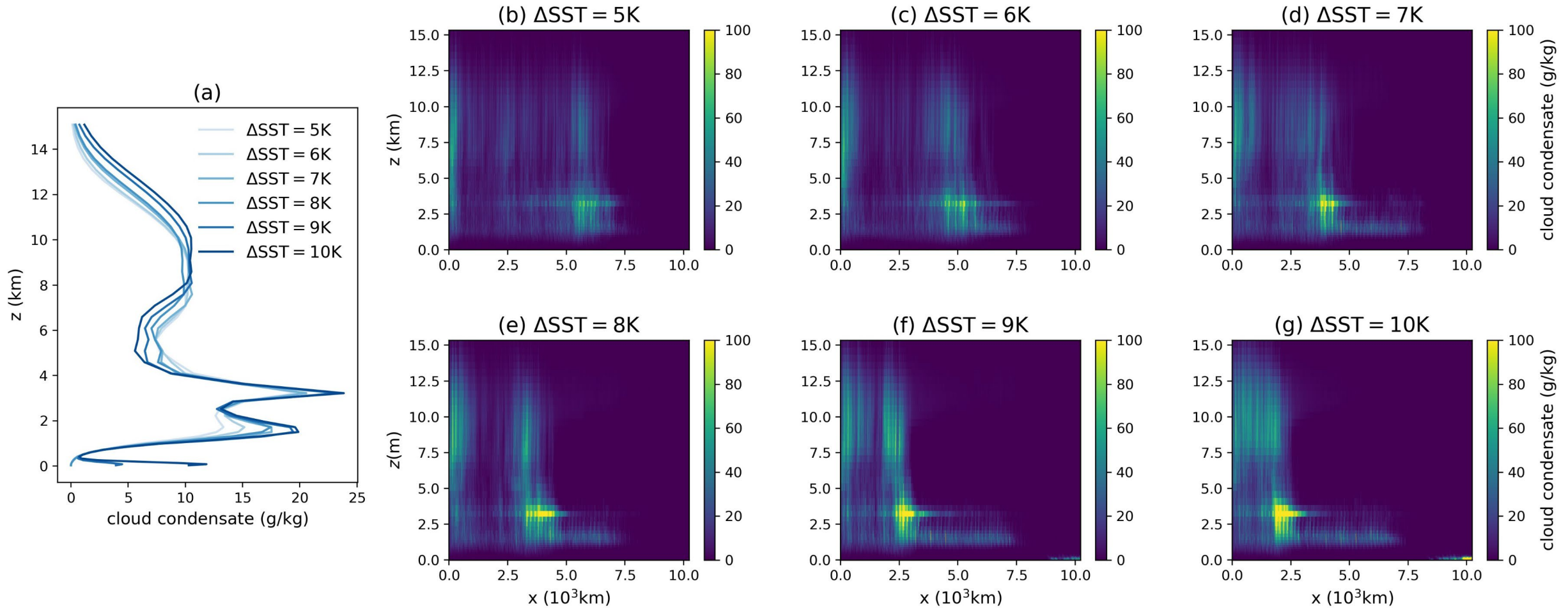
# Stronger convection aggregation $\rightarrow$ larger OLR: mechanism



RH( $x, z$ ) for different  $\Delta$ SST

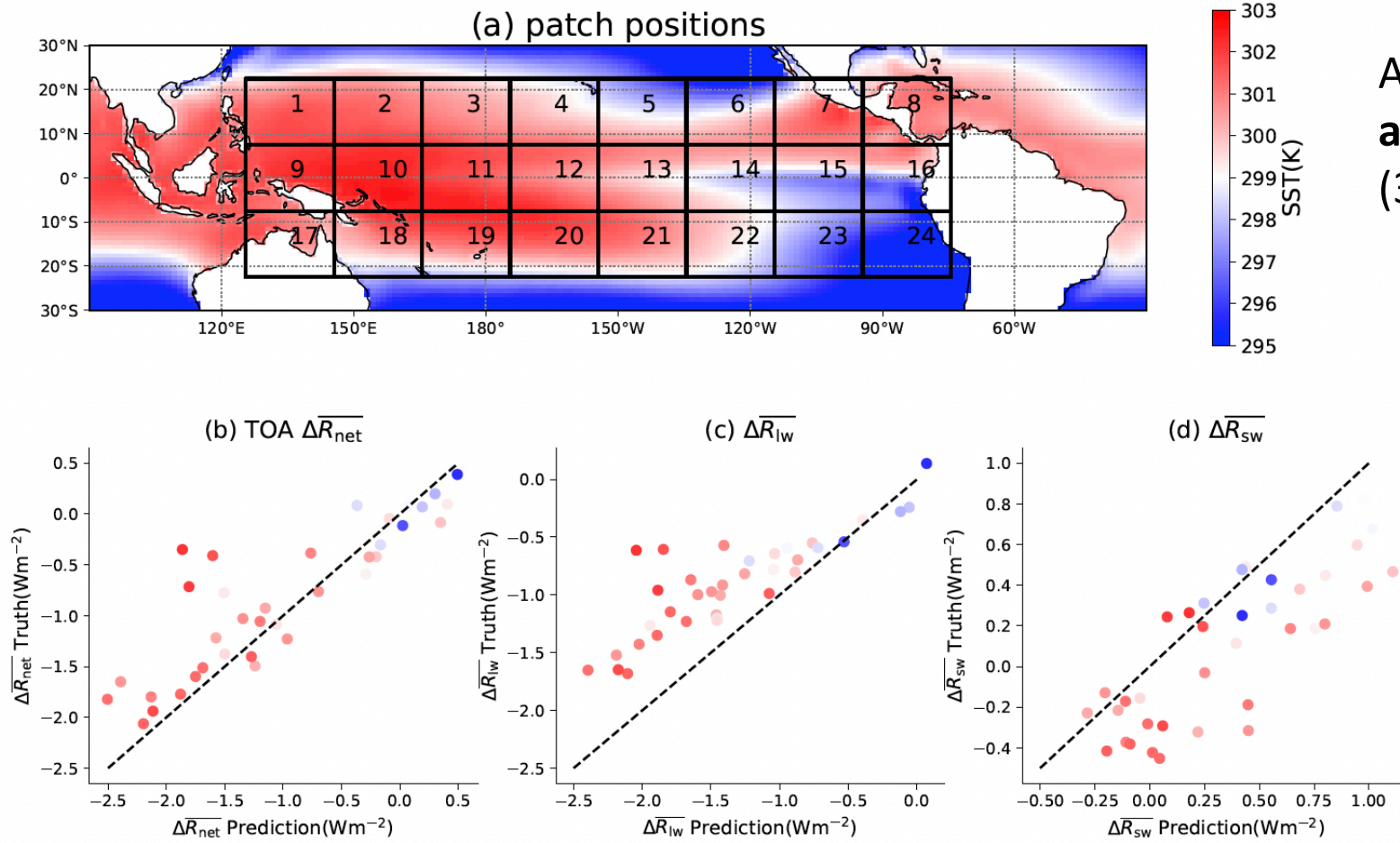


# Stronger convection aggregation $\rightarrow$ larger OLR: mechanism



Cloud condensate mixing ratio for different  $\Delta$ SST

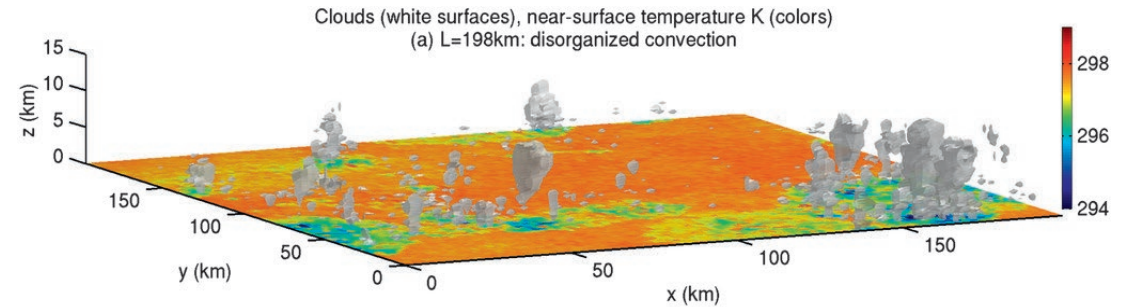
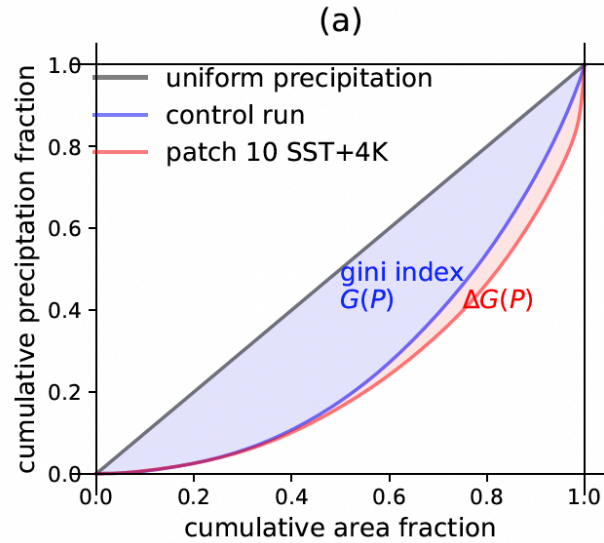
# Linear sum overestimates TOA radiation responses



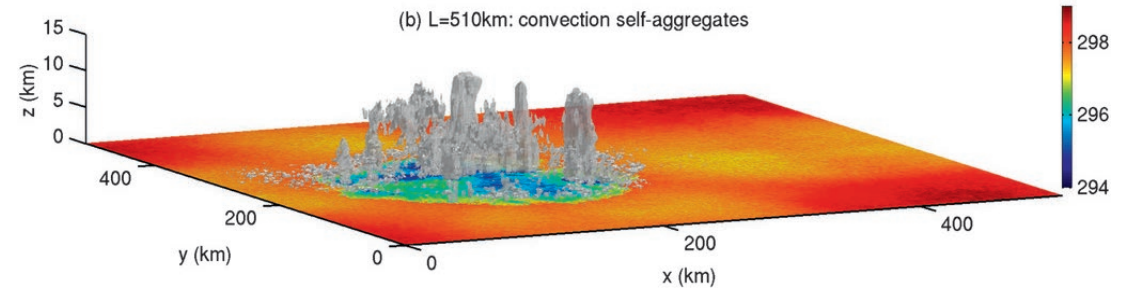
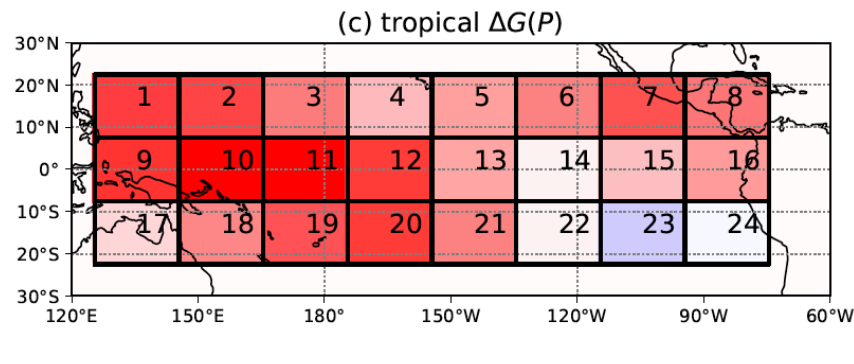
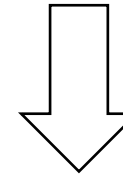
AM4, SST+4K perturbation in **two adjacent** tropical Pacific patches (37 combinations in total)

$$\overline{\Delta R_{net}} = \overline{\Delta R_{lw}} + \overline{\Delta R_{sw}}$$

# Gini index measures large-scale convection aggregation strength



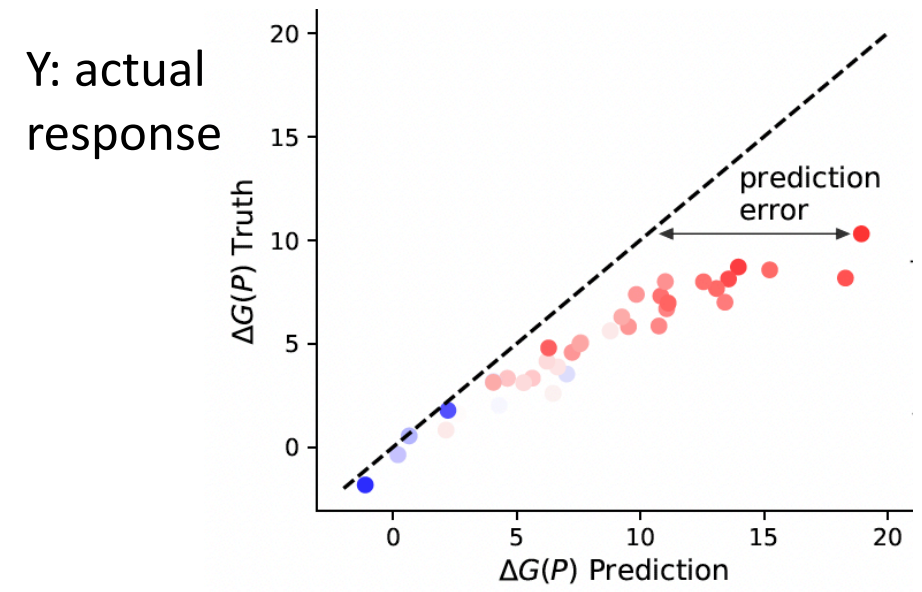
$$\Delta G(P) > 0$$



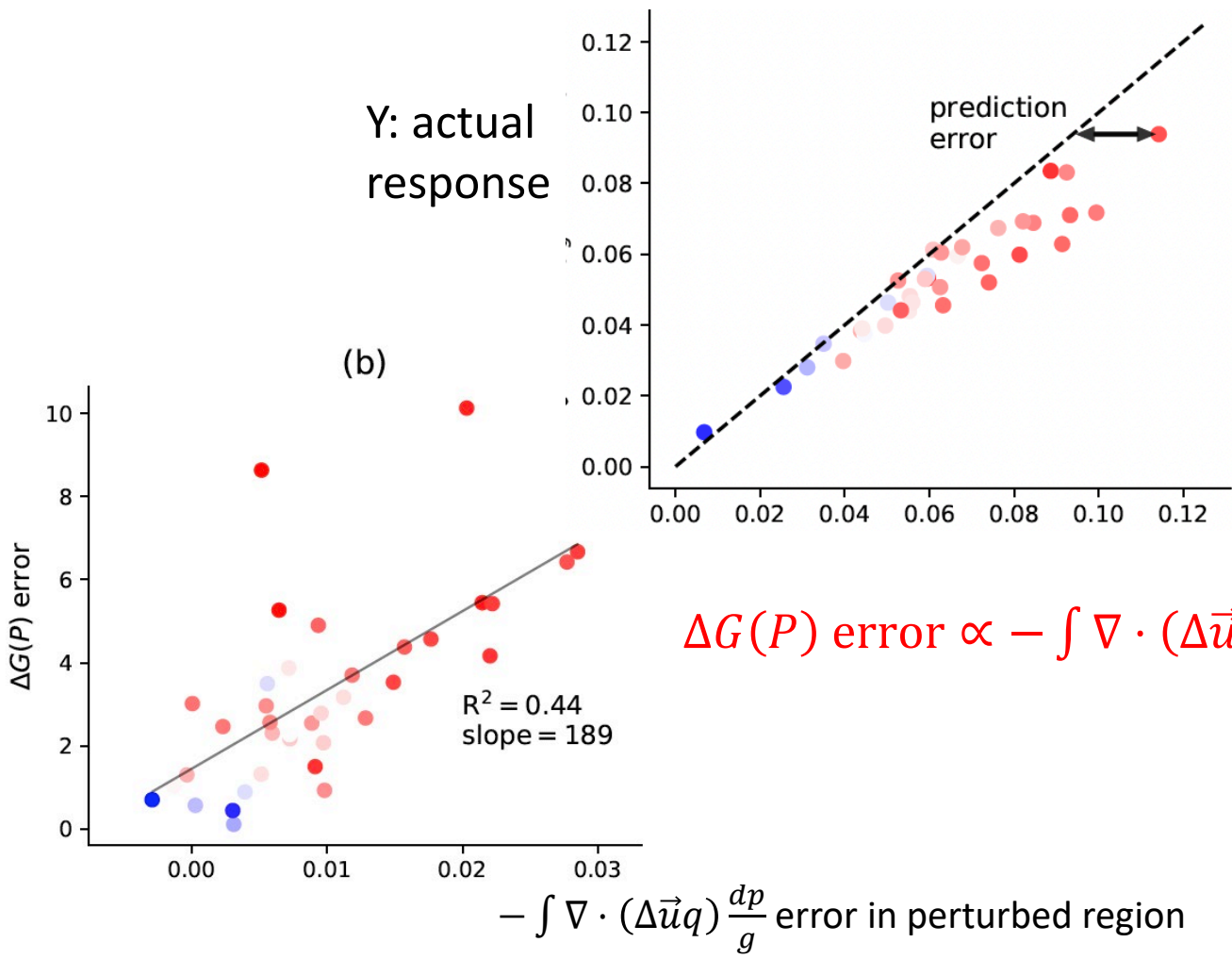
single patch SST+4K  $\Delta G(P)$  responses

Muller & Held, 2012

# convection aggregation overestimation ← circulation overestimation



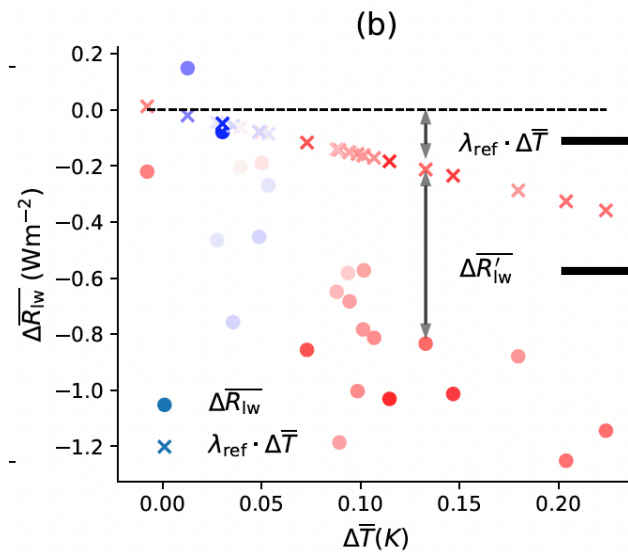
two adjacent patches SST+4K



$$\Delta G(P) \text{ error} \propto - \int \nabla \cdot (\Delta \vec{u} q) \frac{dp}{g} \text{ error}$$

$$- \int \nabla \cdot (\Delta \vec{u} q) \frac{dp}{g} \text{ error in perturbed region}$$

# Longwave radiation response attributed to convection aggregation

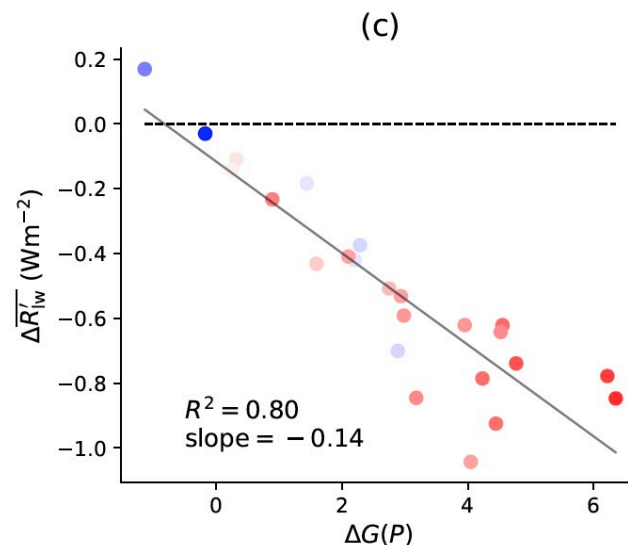


$\lambda_{\text{ref}} \cdot \Delta \bar{T}$  due to mean surface warming

$\Delta \overline{R'_{lw}}$  due to convection aggregation (dominant)

$$\Delta \overline{R_{lw}} = \lambda_{\text{ref}} \cdot \Delta \bar{T} + \Delta \overline{R'_{lw}}.$$

(single patch SST+4K)



$$\Delta \overline{R'_{lw}} \propto \Delta G(P)$$

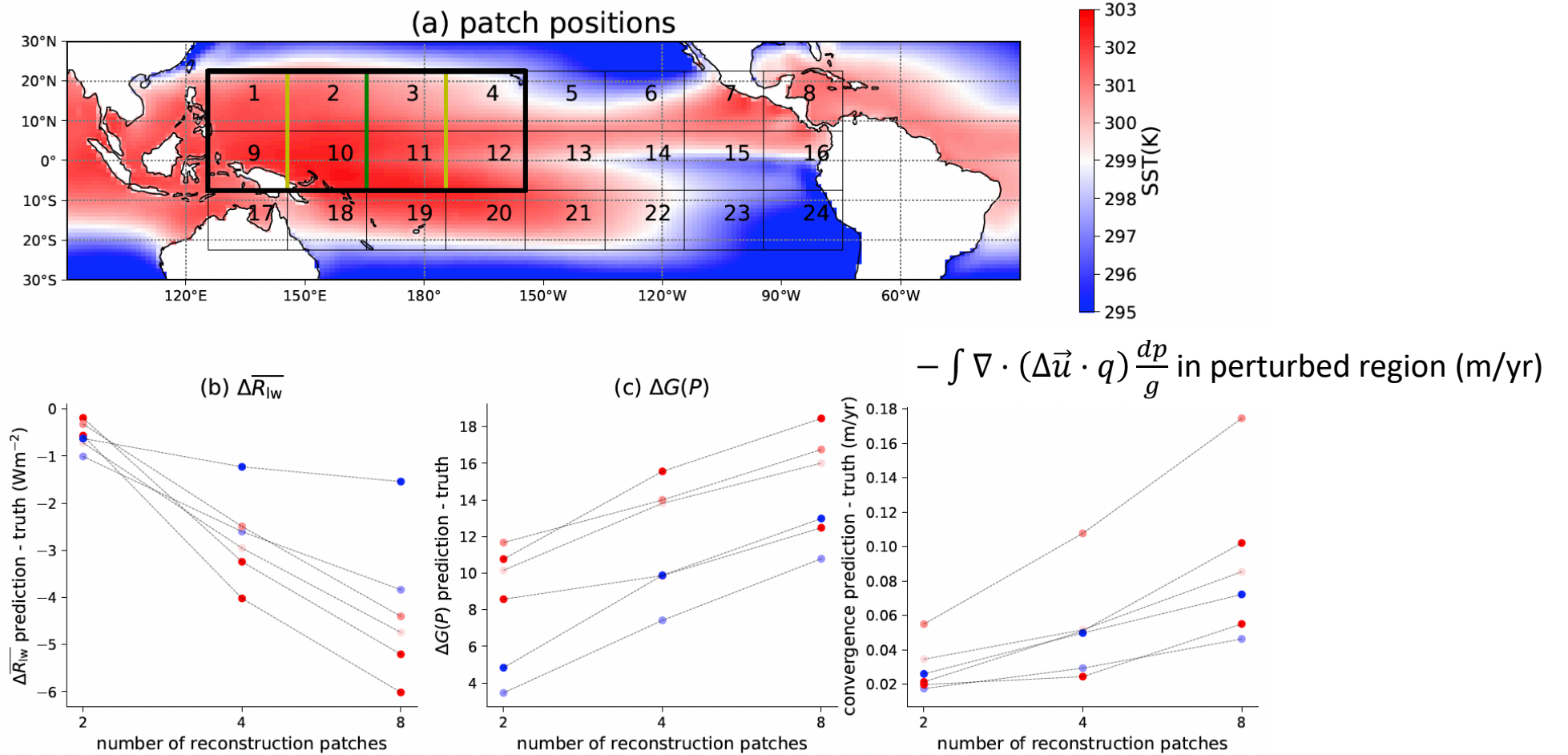
Stronger convection aggregation

→ mid-troposphere drying & high cloud reduction

→ Stronger longwave radiative cooling

Bony 2020, Wing 2020, Zhang 2021

# More (/smaller) patches $\rightarrow$ larger errors

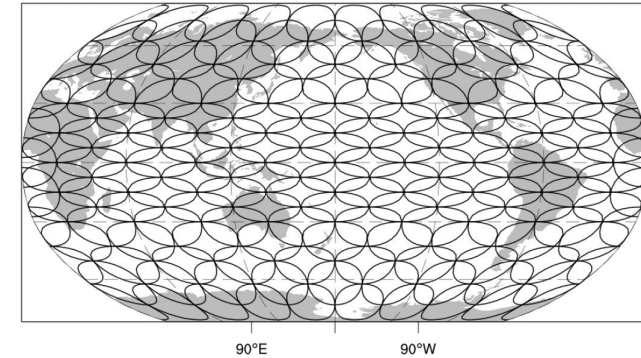
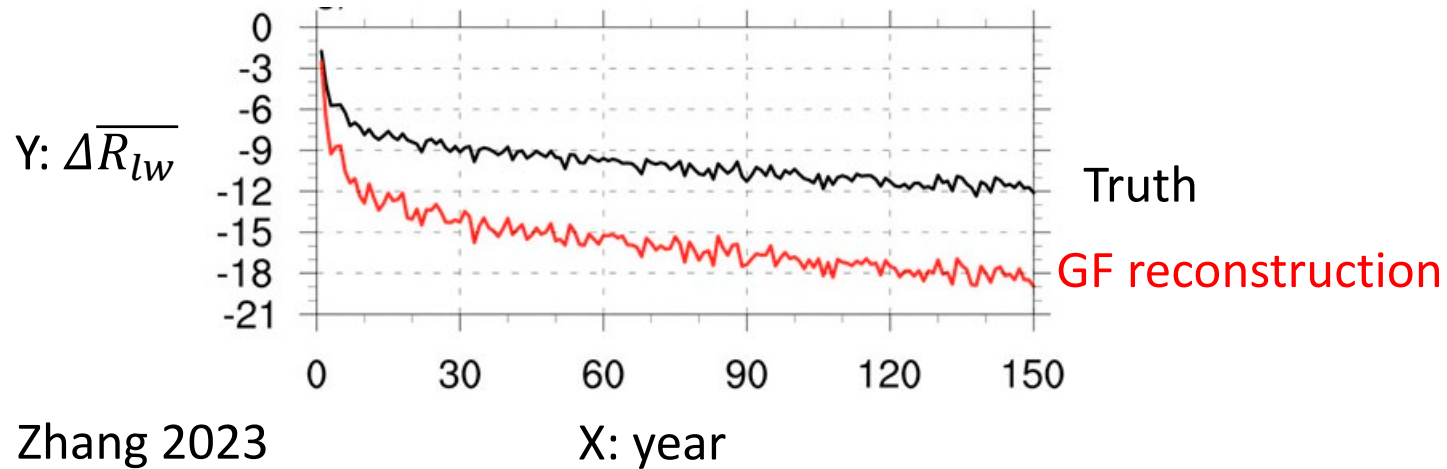


Y: Prediction error

X: Number of reconstruction patches

# The failure of the Green's Functions approach

AGCM  $\leftarrow$  SST warming in 4xCO<sub>2</sub> GCM



Zhang 2023

	AM4 Control	AM4 4×CO <sub>2</sub>	GF reconstruct
tropical $G(P)$	43.9	42.0	74.3
$\overline{MTH}(\%)$	43.1	43.0	34.8
$\overline{HCC}(\%)$	37.8	37.5	30.2

The Green's Functions approach fails in AM4 4xCO<sub>2</sub> radiation reconstruction due to the **overestimation of convection aggregation**